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# HF and ELF facilities as the educational tools for Geospace research

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# HF facility

- High accuracy Doppler shift measurements (long-term stability  $\sim 10$  ppb per day)
- High sensitivity and performance
- Fully automated and remotely controlled
- Low-cost, software-defined, mounted inside standard PC

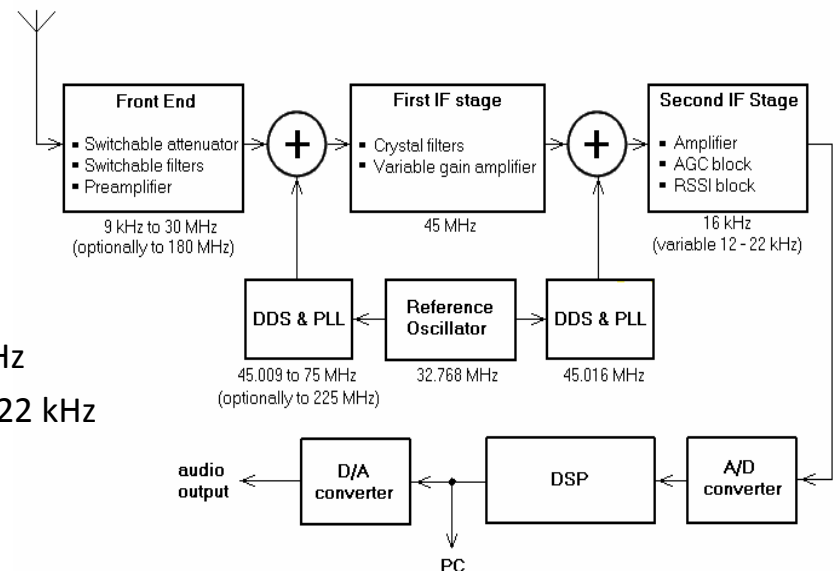
Reference signal source



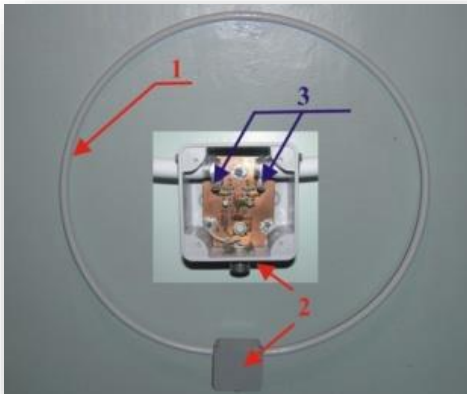
HF Doppler receiver (Winradio G313i) up to 3 units in one PC



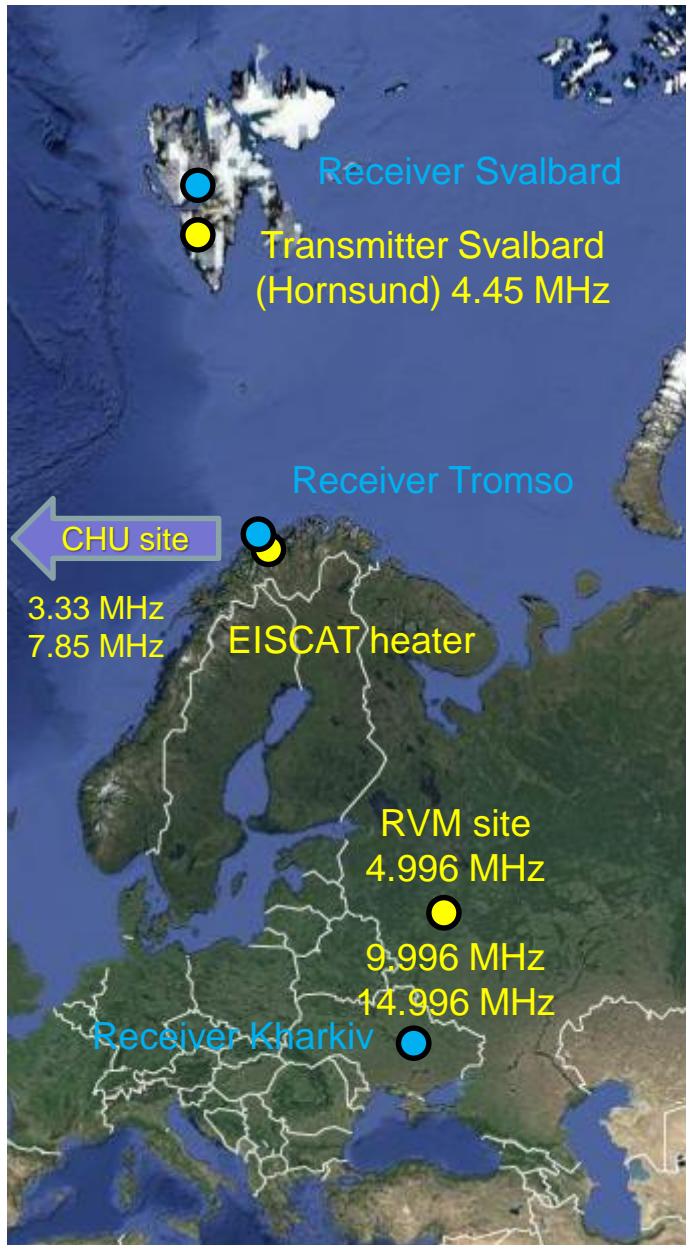
Active loop antenna



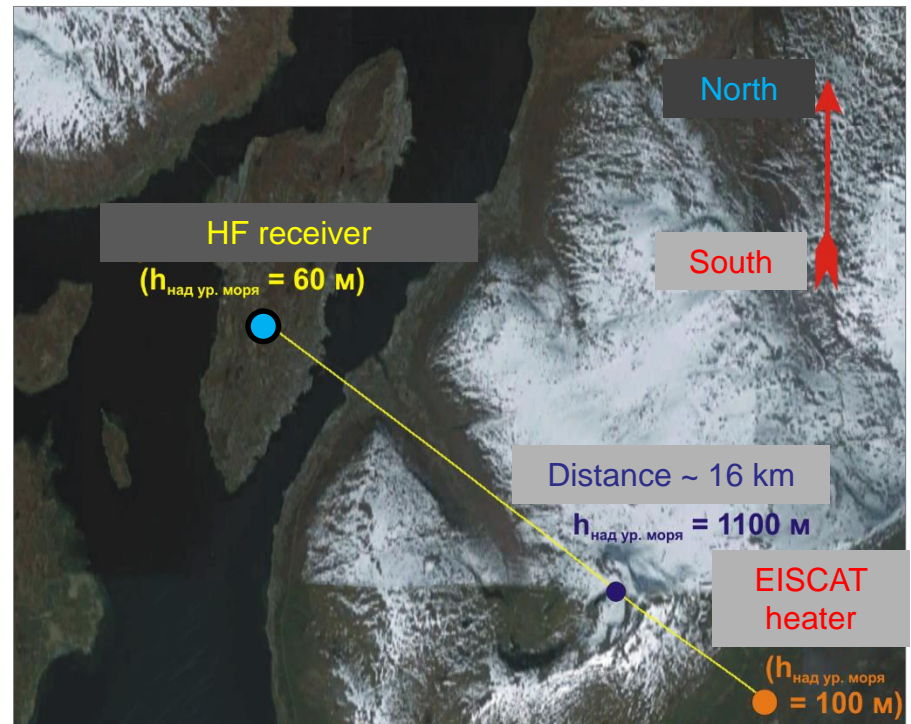
- Frequency Band 3-30 MHz
- Receiver bandwidth 0.5-22 kHz
- Tuning accuracy 1 Hz



# Location

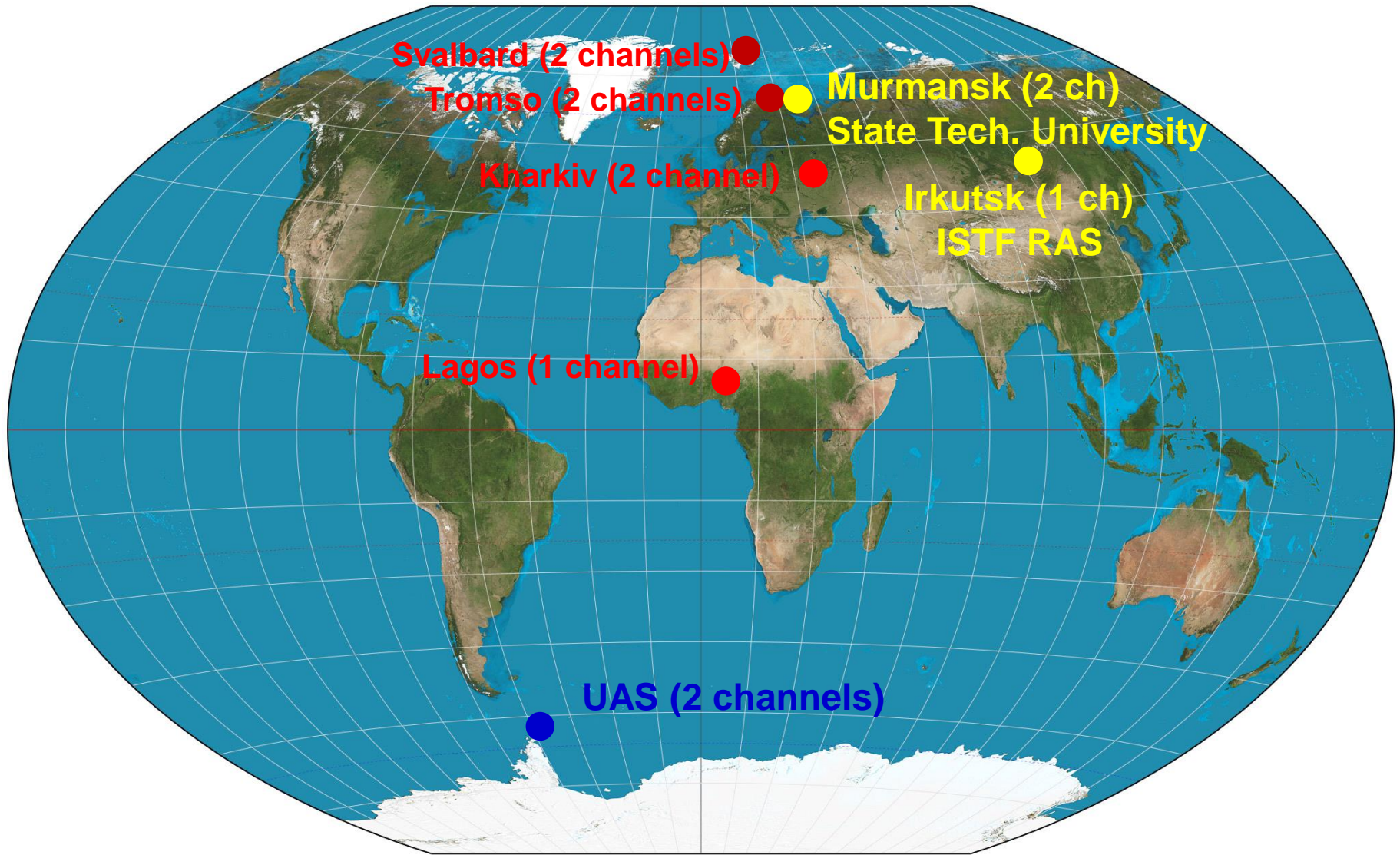


Tromso





# Network of the HF receivers



● Internet-controlled

● Manually-controlled

● Partner instruments

# ELF facility at Svalbard



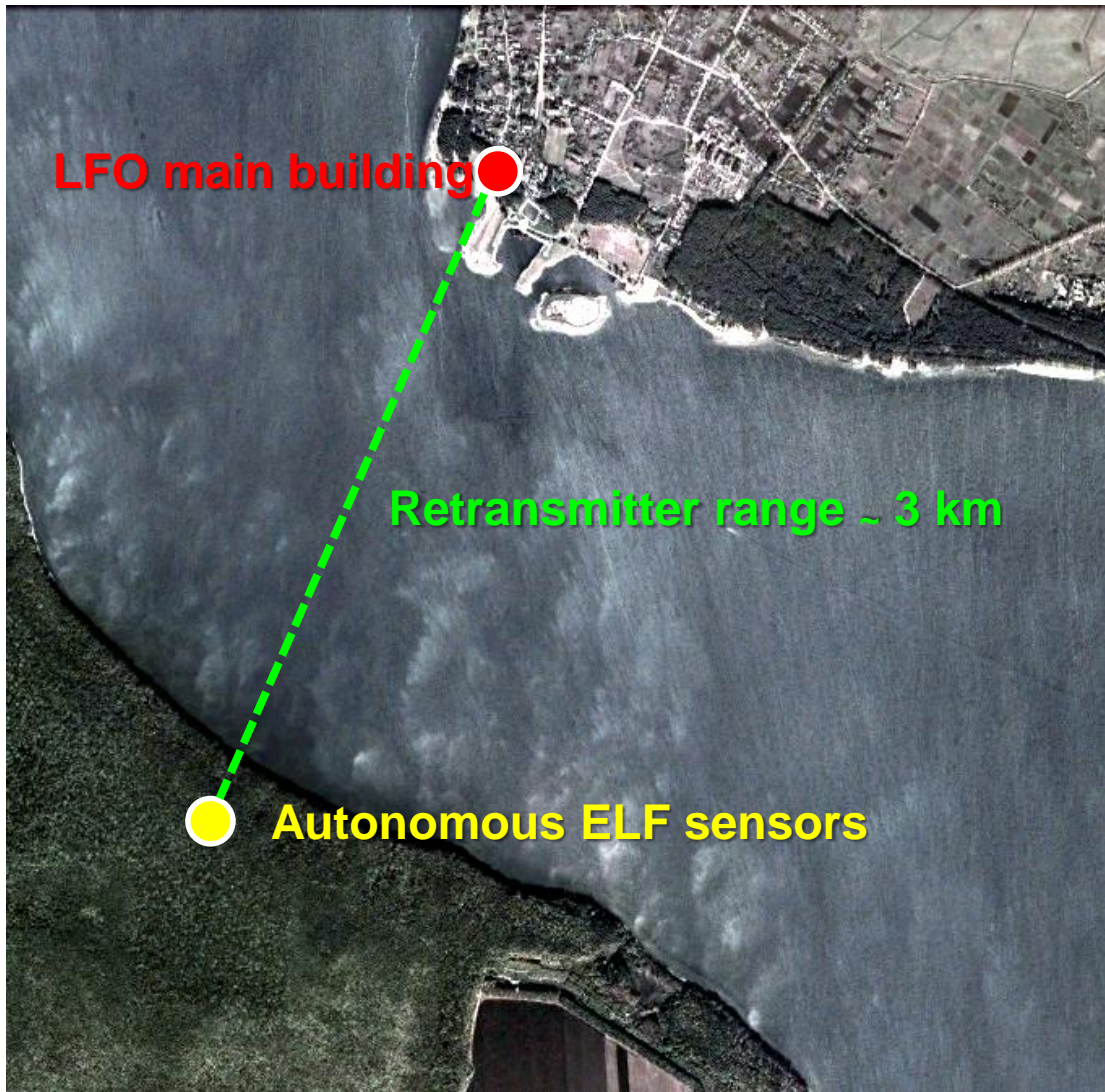
## Lemi-112E

Frequency band (wide)	0.001 - 80 Hz
Sampling rate	254 Hz
Magnetic noise level at 0.001 Hz	$< 200 \text{ pT}/\sqrt{\text{Hz}}$
Magnetic noise level at 0.01 Hz	$< 20 \text{ pT}/\sqrt{\text{Hz}}$
Magnetic noise level at 1 Hz	$< 0.2 \text{ pT}/\sqrt{\text{Hz}}$
Magnetic noise level at 10 Hz	$< 0.03 \text{ pT}/\sqrt{\text{Hz}}$



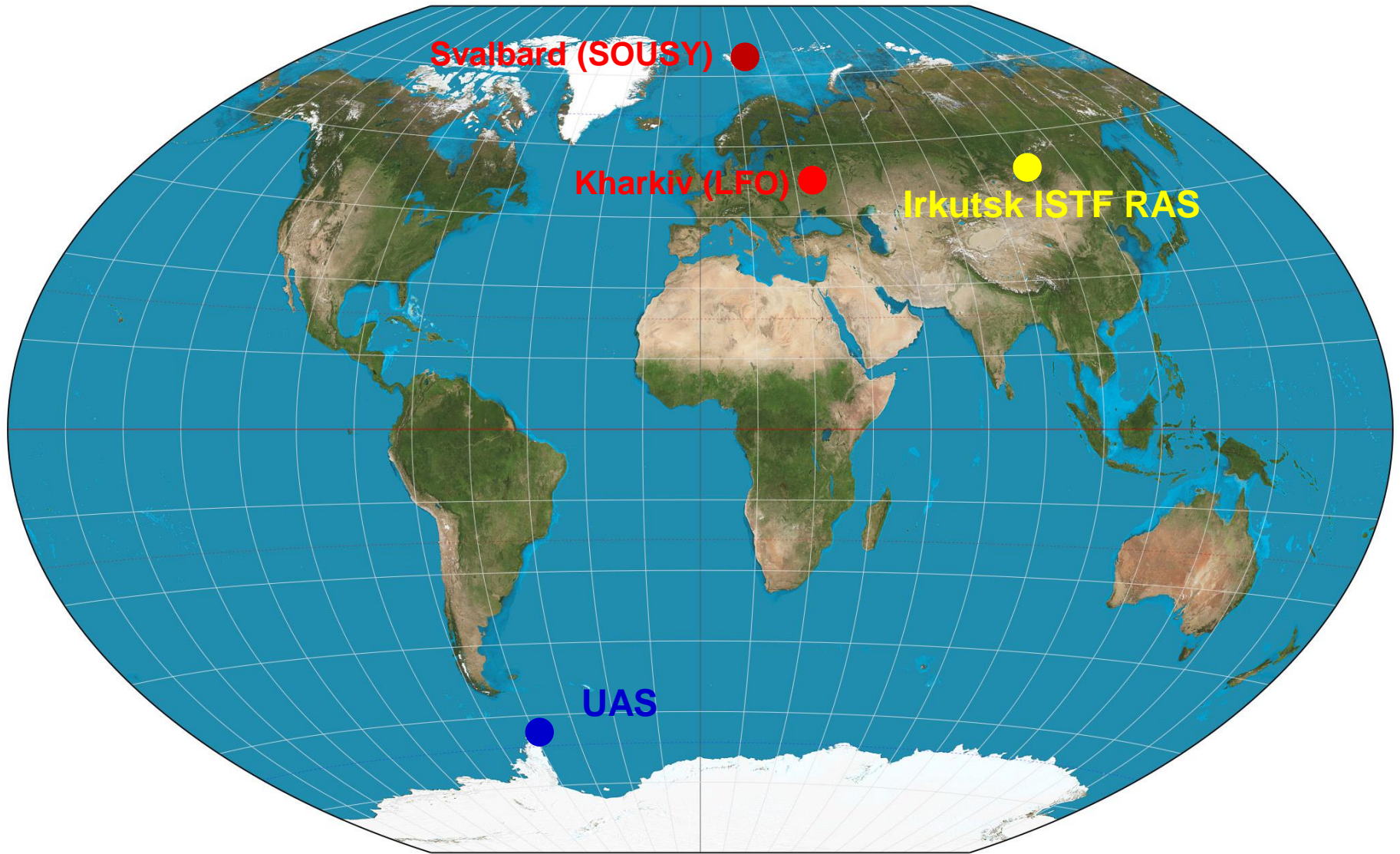


# ELF facility near Kharkiv



Frequency range	0.5 - 40 Hz
Sampling rate	125 Hz
Retransmitter frequency	433 MHz
Retransmitter power	1 mW
Retransmitter range	≈ 3 km

# Network of the ELF magnetometers



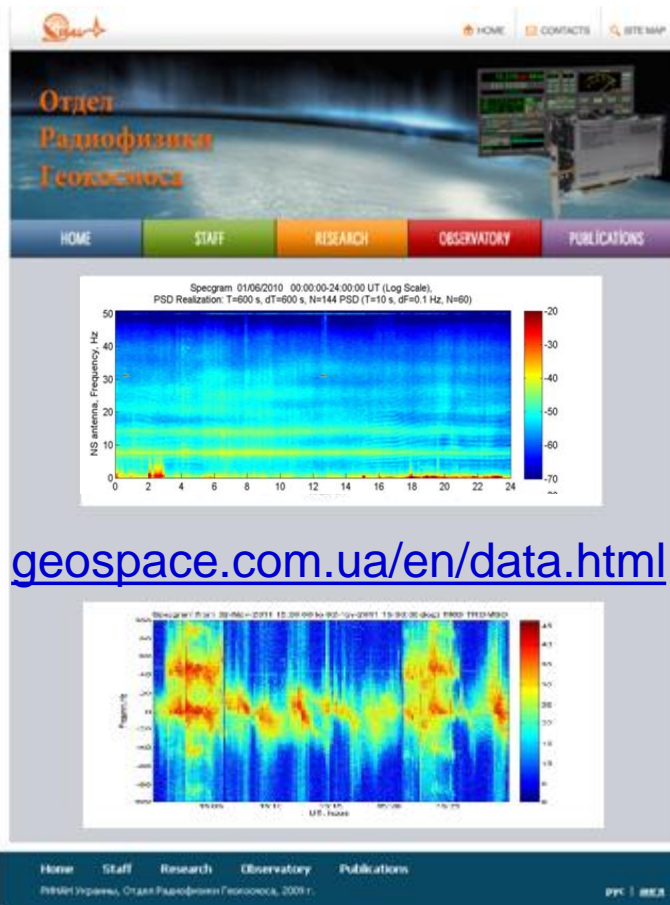
● Internet-controlled

● Manually-controlled

● Partner instruments

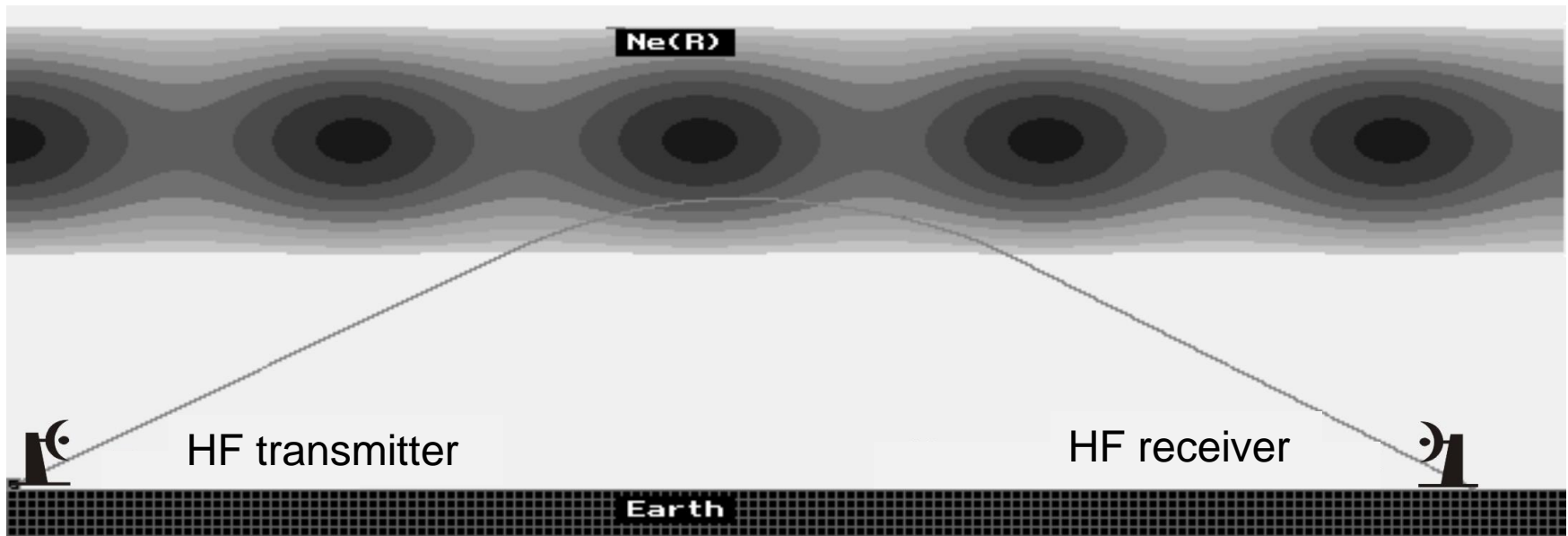


# Remote control and real-time Internet access to data, data visualizing





# Layout of the HF observations



$$\Delta f = -\frac{f}{c} \frac{dP}{dt}, \quad P = \int_S n \cdot dS$$

$\Delta f$  - Doppler frequency shift  
 $P$  - Phase path of the wave in the ionosphere

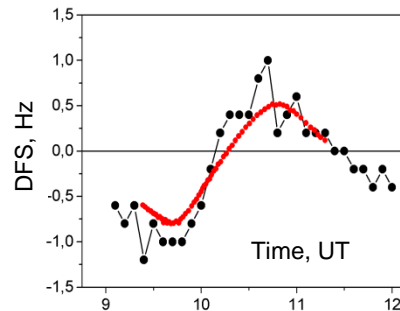
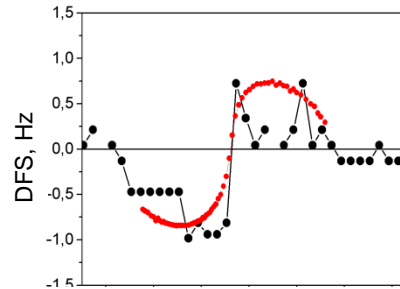
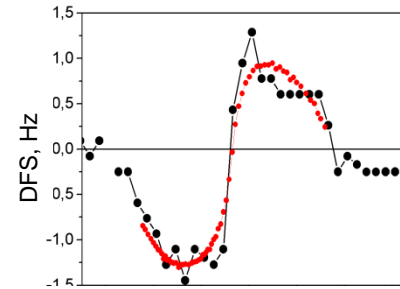
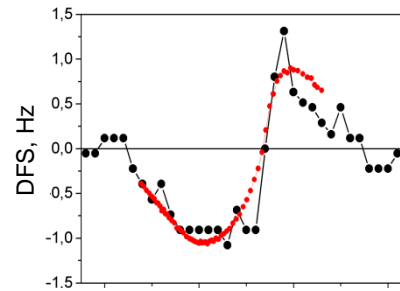
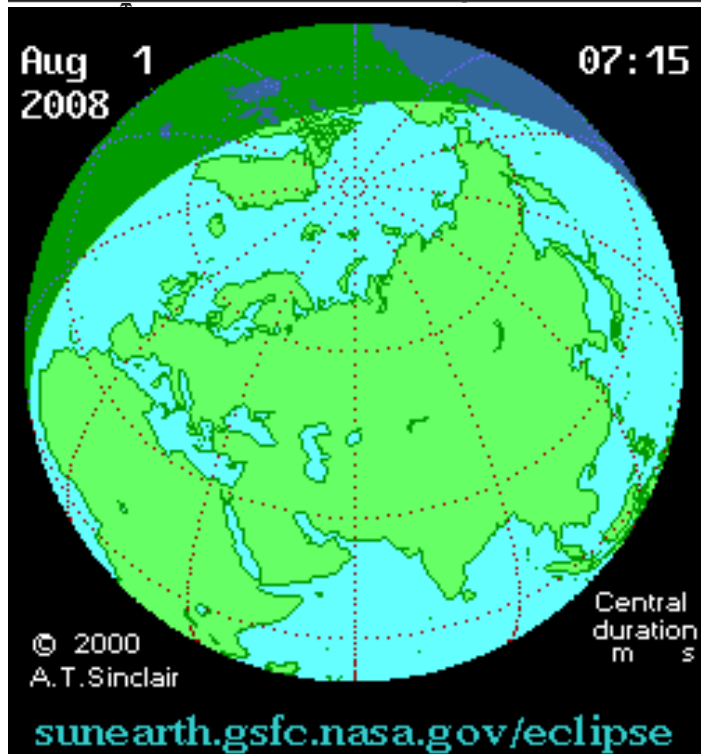
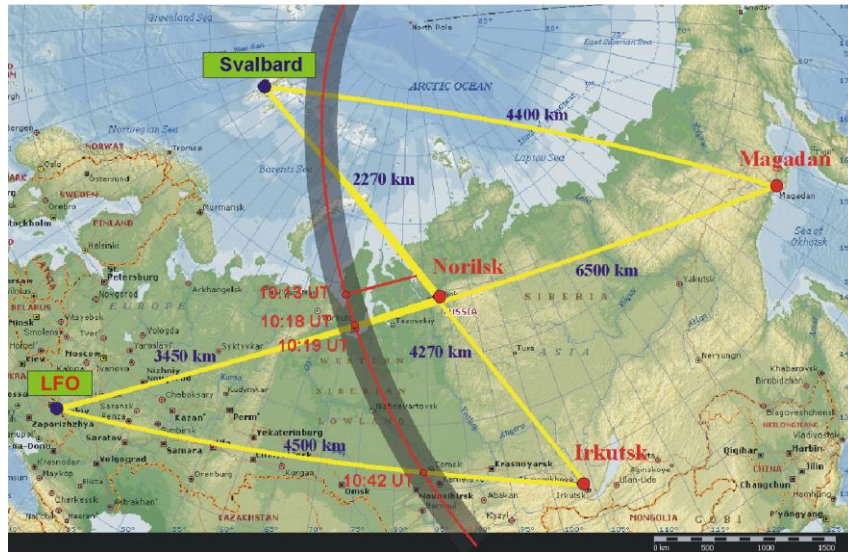
**Phase path  $\Delta P$  varies because of changes of the:**

$n$  - refractive index  
 $S$  - ray trajectory  
 $f$  - signal frequency  
 $c$  - velocity of light

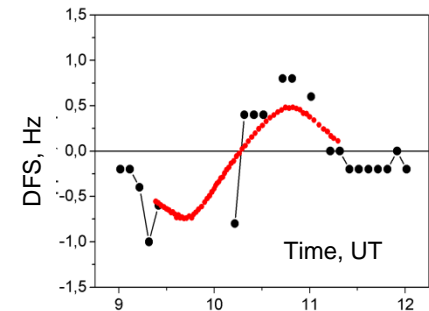
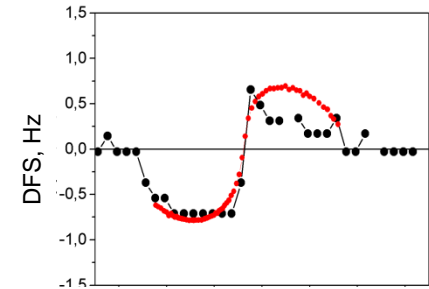
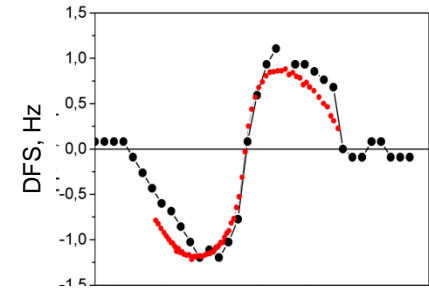
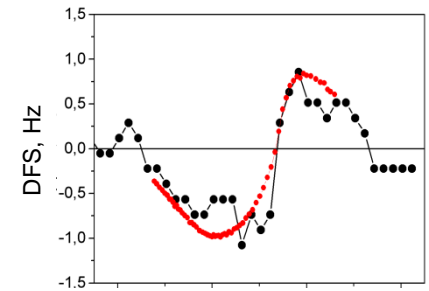
1) Medium  $n(t)$  - then:  $\Delta f \sim \frac{1}{f}$

2) Variations of the ray trajectory  $S(t)$  - then:  $\Delta f \sim f$

# Doppler frequency shift variations during solar eclipse, observed at Svalbard and in Kharkiv



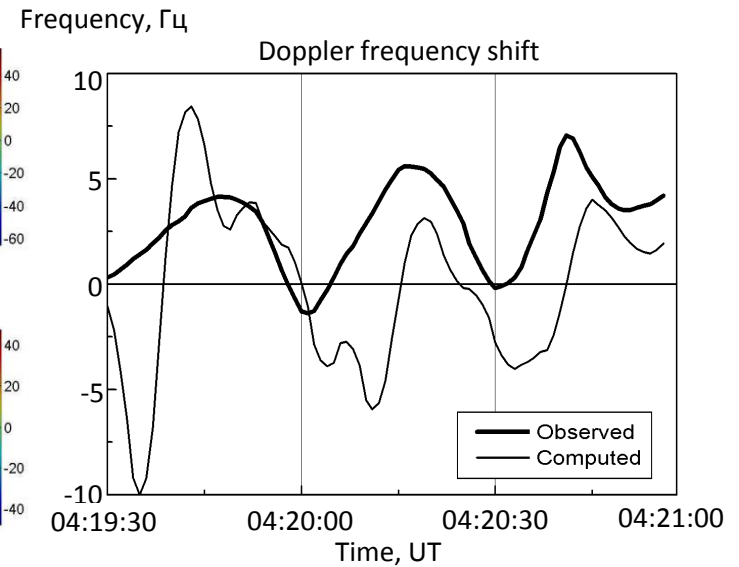
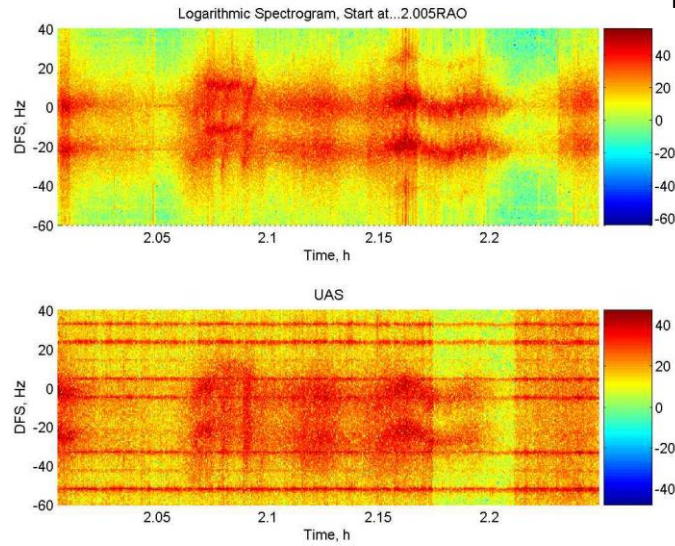
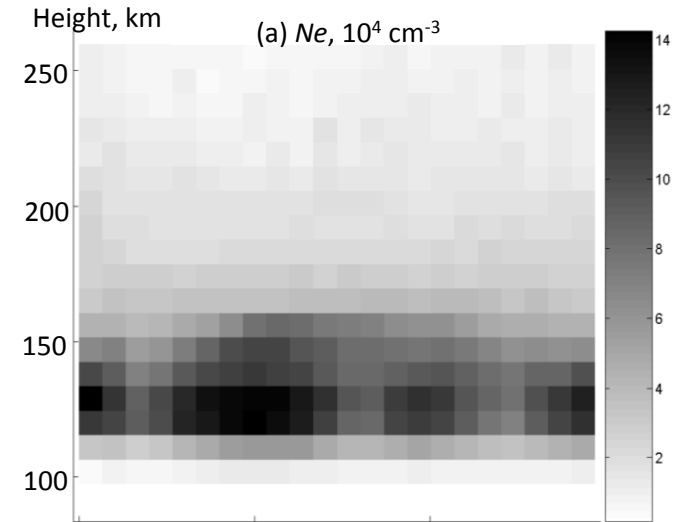
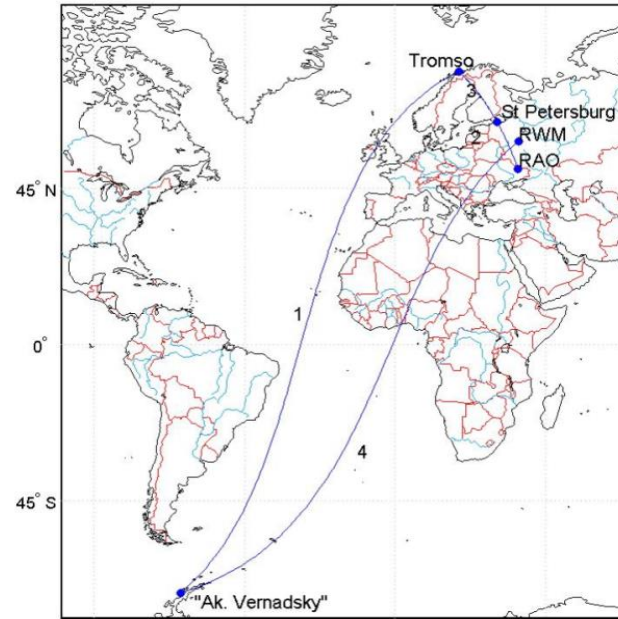
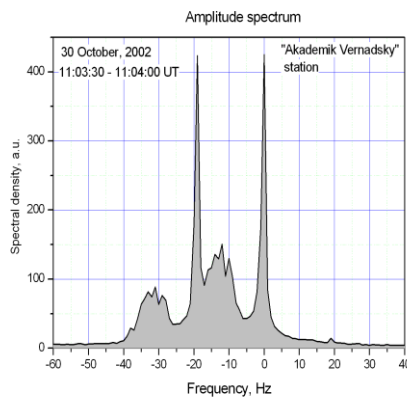
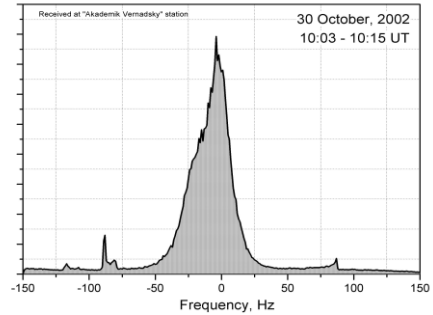
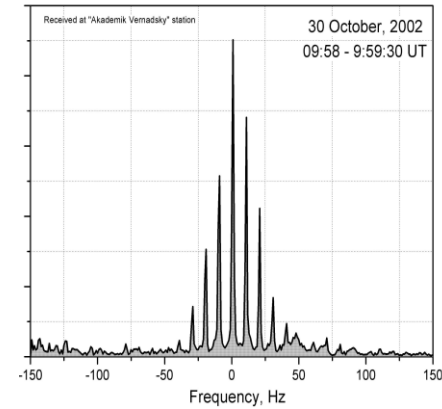
$F = 10999 \kappa \Gamma \varphi$



$F = 12333 \kappa \Gamma \varphi$



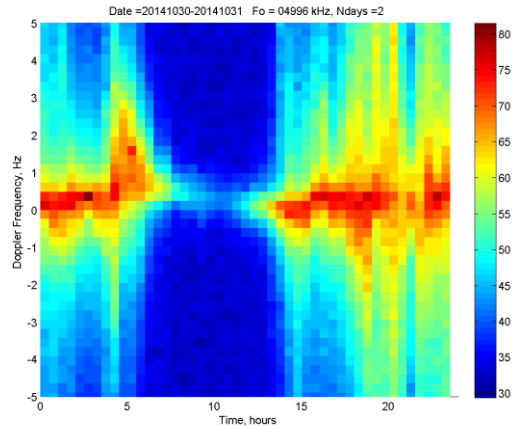
# Self-scattering experiment



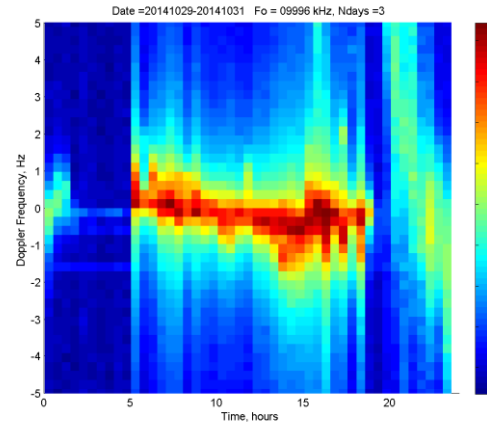
# Monitoring of HF signals of RWM station situated near Moscow (Russia)

RWM -Tromso, 29-31.10.2014

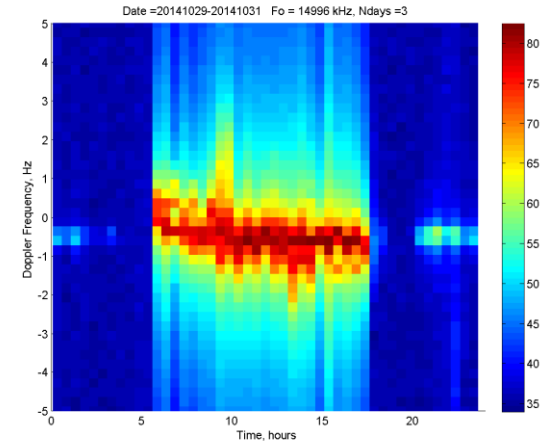
4996 kHz



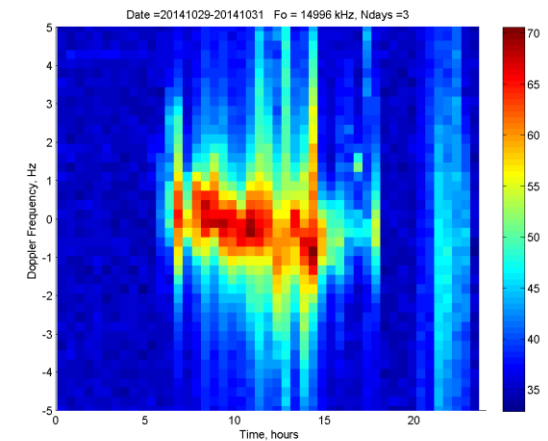
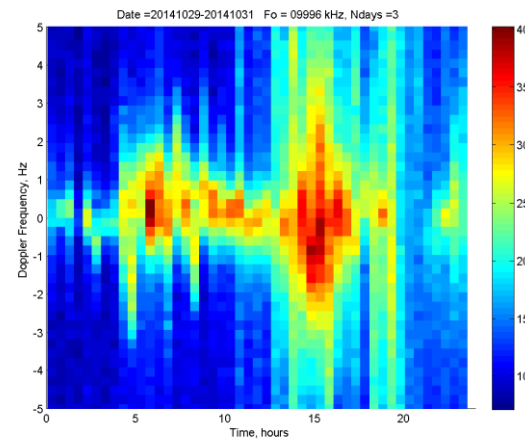
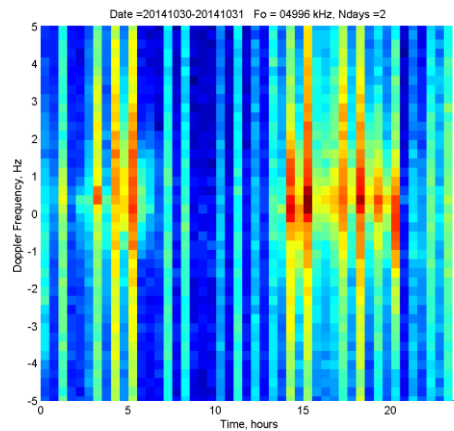
9996 kHz



14996 kHz



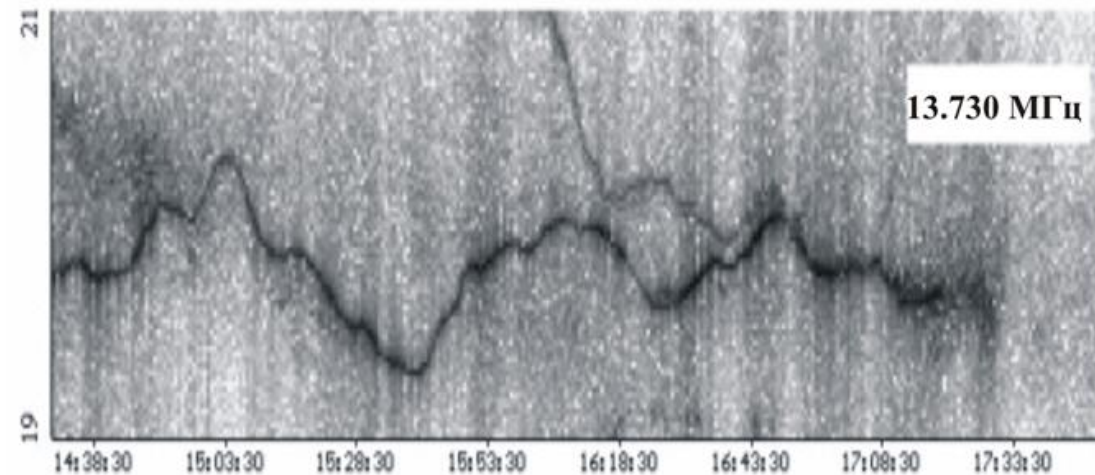
RWM-Svalbard, 29-31.10.2014



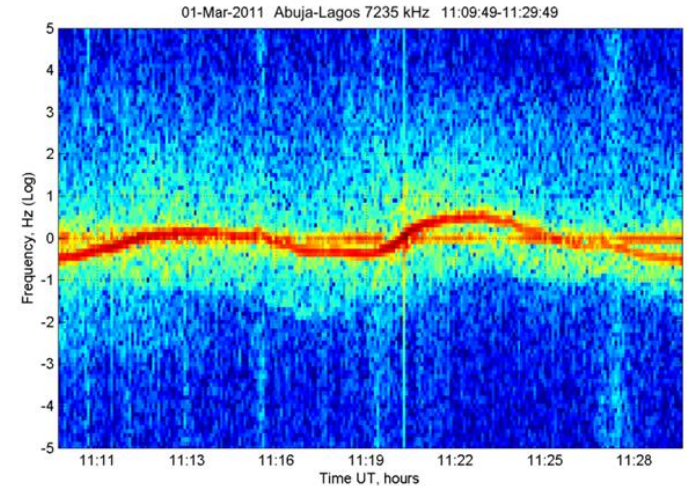


# Doppler shift variations produced by gravity waves and geomagnetic pulsations

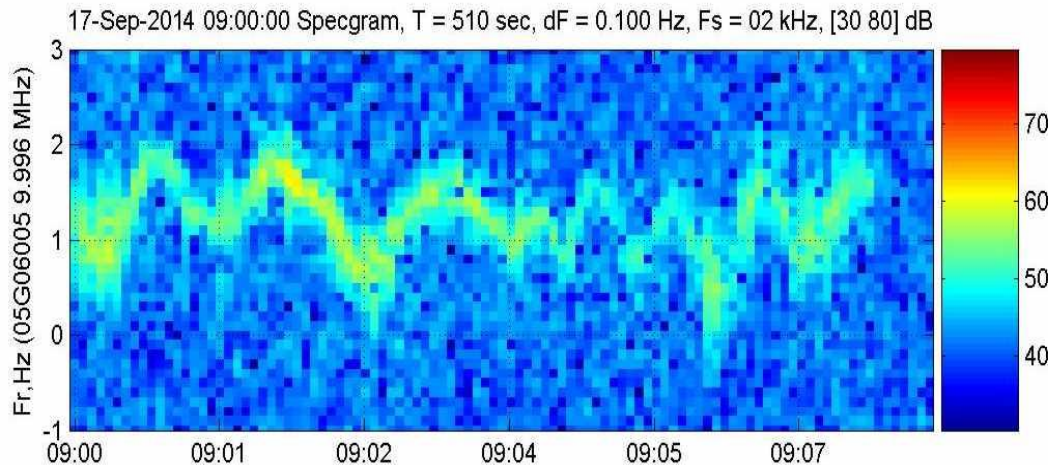
Spectrogram obtained during the transatlantic voyage to Ukrainian Antarctic station



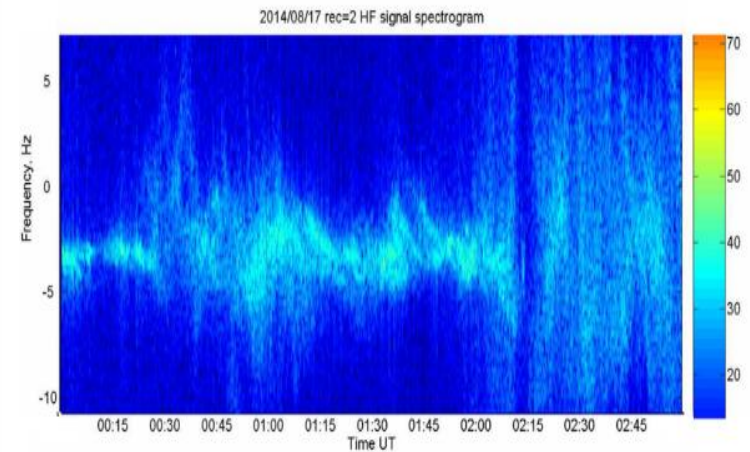
Radio path Abuja-Lagos (Nigeria), D~500 km



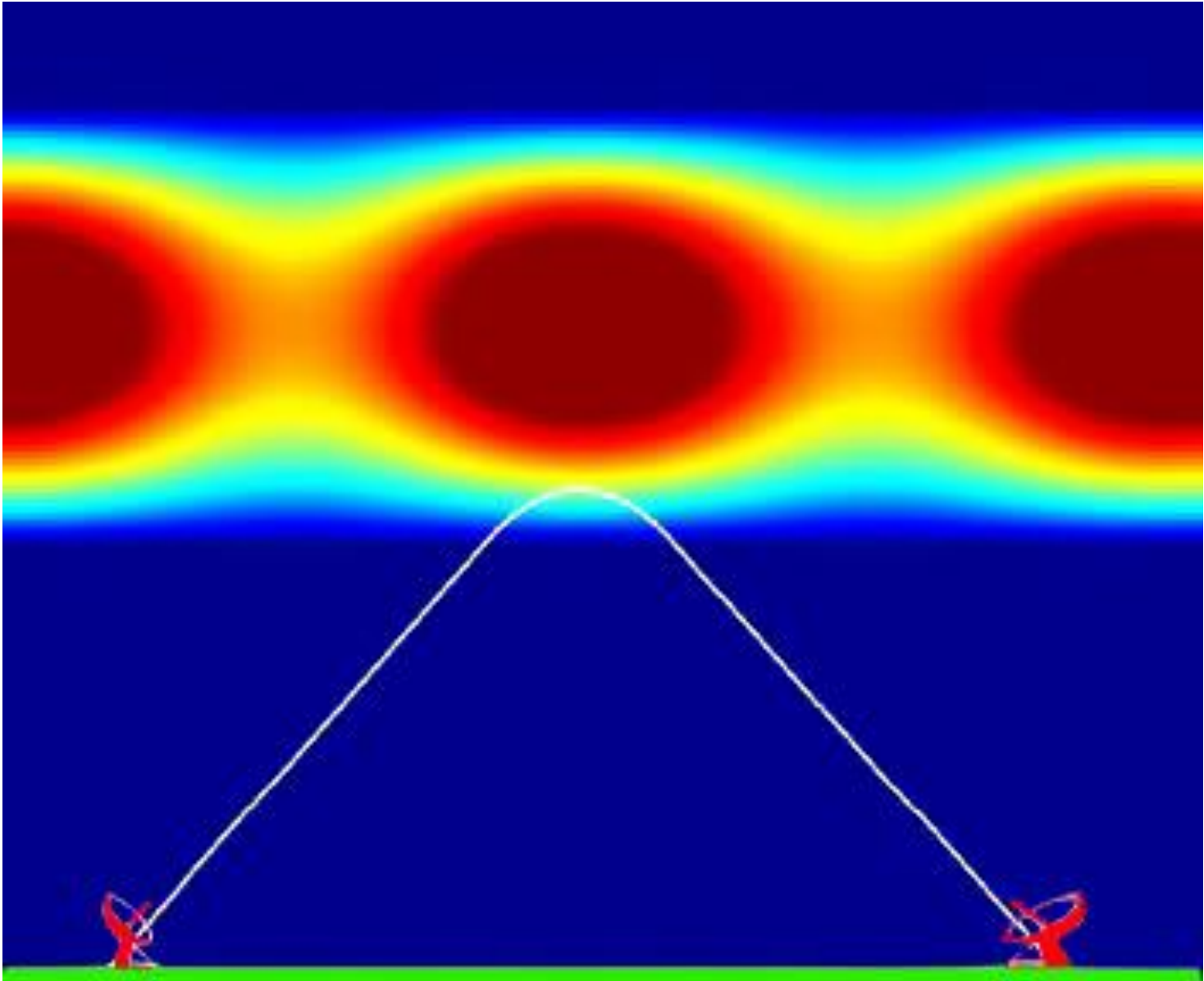
Radio path RWM-Tromso (Norway), D~ 1800 km



Radio path KHO-Hornsund (Svalbard), D~90 km



# Traveling ionospheric disturbances (model)







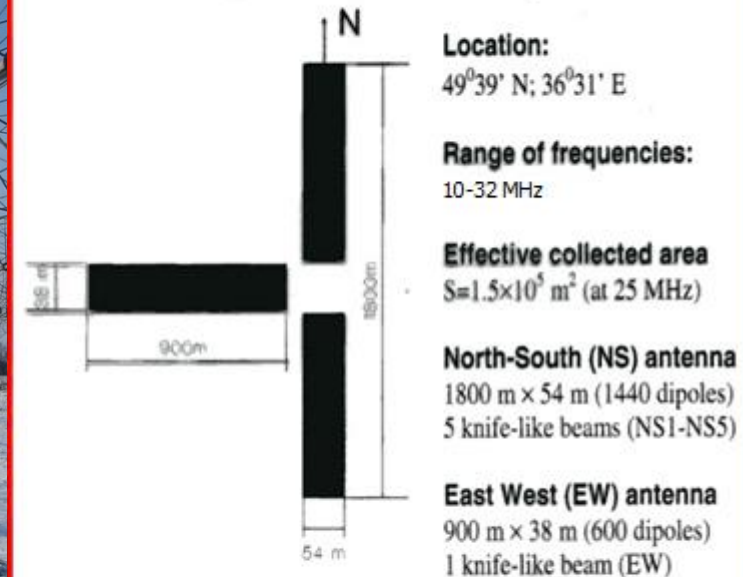
al\_geronimo.livejournal.com



al\_geronimo.livejournal.com

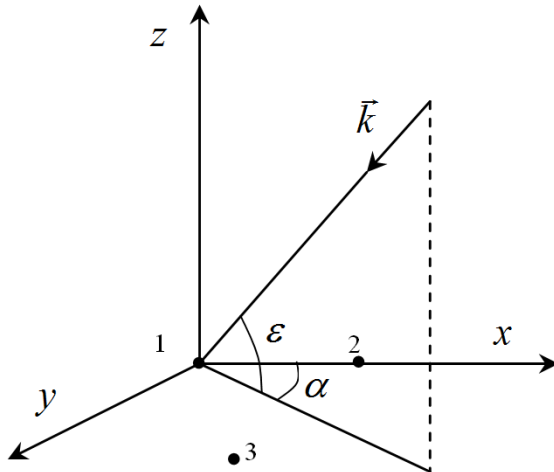


## UTR-2 radio telescope



- All antennas are independently electrically steerable with a discrete step  $0.2^{\circ}$
- The characteristic pattern widths of the UTR-2 beams when phased toward heated volume  $\Delta\epsilon \sim 1^{\circ}; \Delta A \sim 1.5^{\circ}$

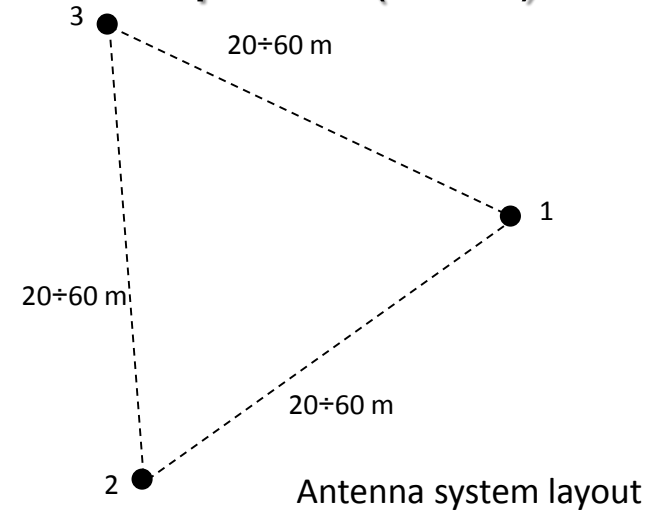
# Frequency and Angular Sounding of the Ionosphere (FASI)



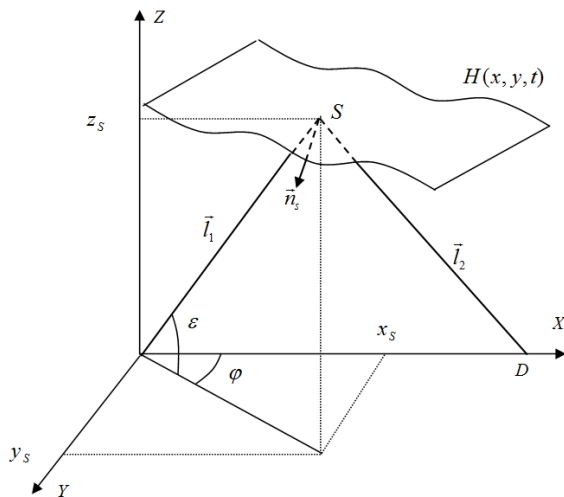
$$\tan \alpha = \frac{x_2 \Delta \varphi_{3-1} - x_1 \Delta \varphi_{2-1}}{y_3 \Delta \varphi_{2-1}}$$

$$\cos \varepsilon = \frac{\Delta \varphi_{3-1}}{k(x_3 \cos \alpha + y_3 \sin \alpha)}$$

where  $\Delta \varphi_{3-1}, \Delta \varphi_{2-1}$  are phase differences and  $k$  is wave vector.



## Model of perfectly reflecting ionospheric surface



$$H(x, y, t) = H_o [1 + h(x, y, t)] \quad |h| \ll 1$$

$$l_1 + l_2 = l \equiv c \cdot \tau_d \quad \tan \alpha_x = \frac{\partial H}{\partial x} \equiv \gamma_x \quad \tan \alpha_y = \frac{\partial H}{\partial y} \equiv \gamma_y$$

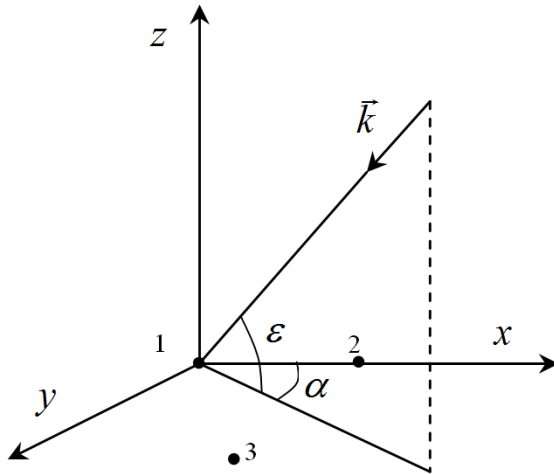
$$|\gamma_x|, |\gamma_y| \ll 1$$

$$\frac{l \cdot \cos \varepsilon \cdot \cos \varphi - D}{\sqrt{D^2 + l^2 - 2 \cdot D \cdot l \cdot \cos \varepsilon \cdot \cos \varphi}} = -\gamma_x \quad \frac{l \cdot \cos \varepsilon \cdot \sin \varphi}{\sqrt{D^2 + l^2 - 2 \cdot D \cdot l \cdot \cos \varepsilon \cdot \cos \varphi}} = -\gamma_y$$

$$F_D = -\frac{1}{\lambda} \frac{dl}{dt} \quad X_s = -\gamma_x \quad Y_s = -\gamma_y \quad F_D = -\frac{4H_0^2}{\lambda \cdot l_0} \cdot \frac{dh}{dt}$$



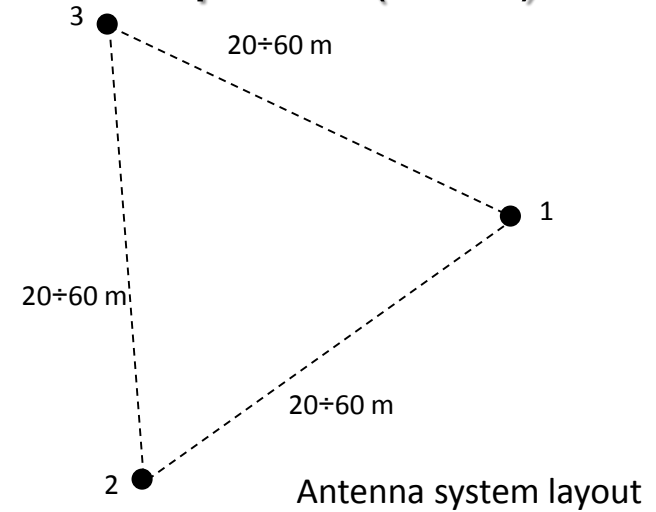
# Frequency and Angular Sounding of the Ionosphere (FASI)



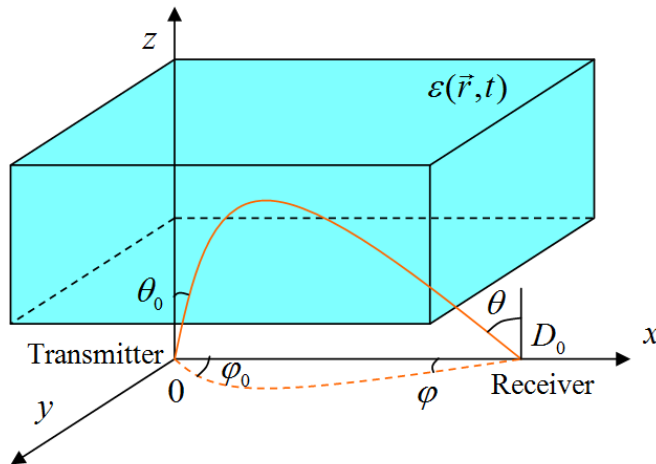
$$\tan \alpha = \frac{x_2 \Delta \varphi_{3-1} - x_1 \Delta \varphi_{2-1}}{y_3 \Delta \varphi_{2-1}}$$

$$\cos \varepsilon = \frac{\Delta \varphi_{3-1}}{k(x_3 \cos \alpha + y_3 \sin \alpha)}$$

where  $\Delta \varphi_{3-1}, \Delta \varphi_{2-1}$  are phase differences and  $k$  is wave vector.



## Model of density waves propagating in a ionospheric layer



$$\varepsilon(\vec{r}, t) = \varepsilon_0(z) + \varepsilon_1(\vec{r}, t) \quad |\max \varepsilon_1| \ll \varepsilon_0$$

$$\varepsilon_0(z) = 1 - \alpha^2 \cdot F(z) \quad \Phi(z) \quad \text{height profile of the disturbances amplitude}$$

$$\varepsilon_1(\vec{r}, t) = \Phi(z) \cdot \nu(\vec{r}, t) \quad \nu(\vec{r}, t) \quad \text{3D disturbance waveform}$$

$$\alpha = f_{cr} / f_0 \quad L = L_0 + L_1 \quad L_0 = \int_0^{D_0 / \sin \theta_0} \varepsilon_0(z(\rho)) d\rho \quad L_1 = \frac{1}{2} \int_0^{D_0 / \sin \theta_0} \varepsilon_1(\vec{r}(\rho), t) d\rho$$

$$F_D = -\frac{1}{\lambda} \frac{\partial L_1}{\partial t}$$

$$\Delta \varphi \approx -\frac{1}{D_0 \sin \theta_0} \frac{\partial L_1}{\partial \varphi_0}$$

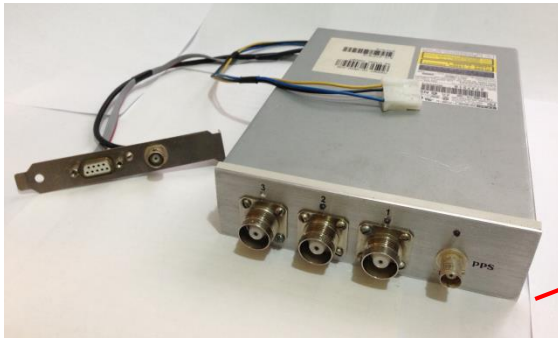
$$\Delta \theta \approx \frac{1}{\cos \theta_0 (\partial D / \partial \theta_0)} \frac{\partial L_1}{\partial \theta_0}$$

# Low-cost FASI complex based on Winradio G3131i and antenna switch

Ovenized crystal oscillator



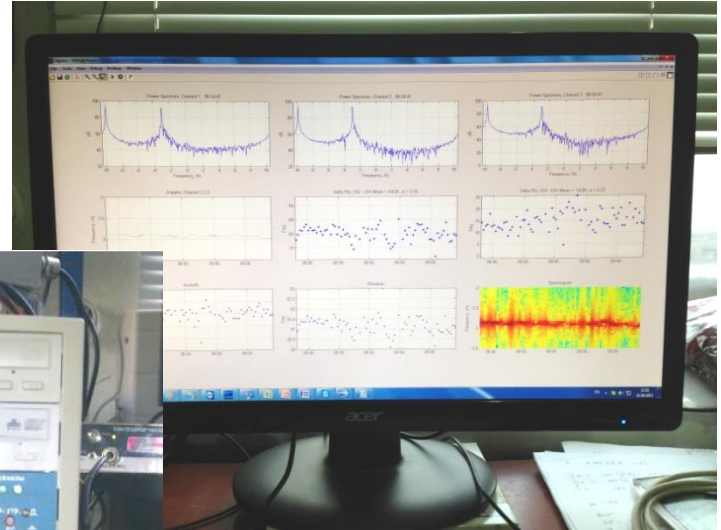
Antenna switch



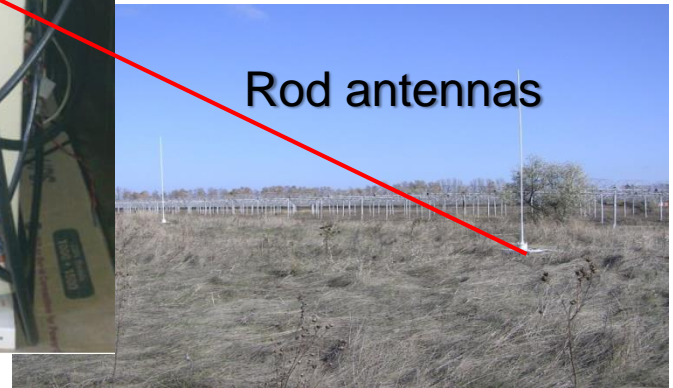
HF receiver



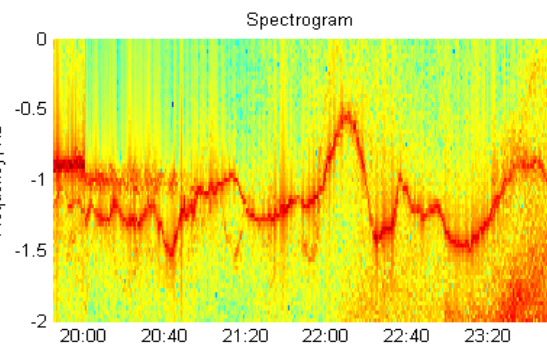
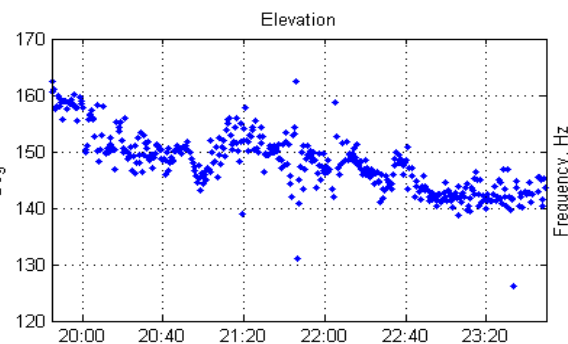
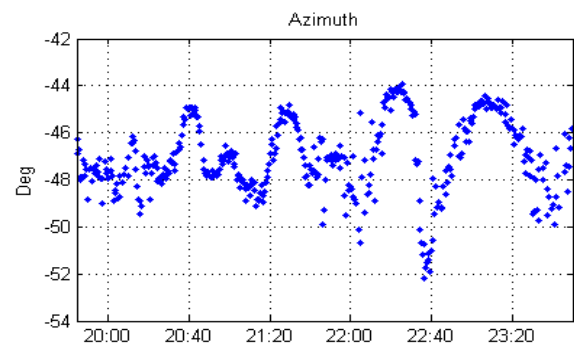
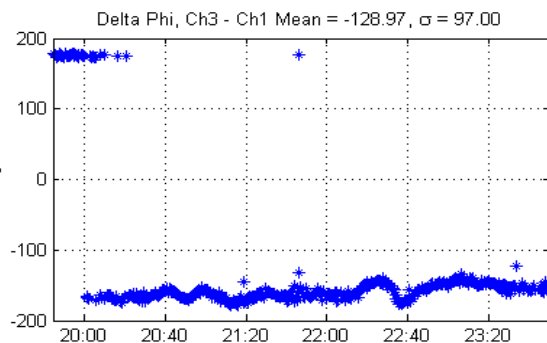
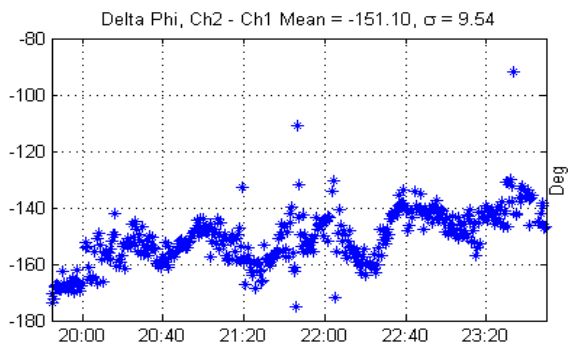
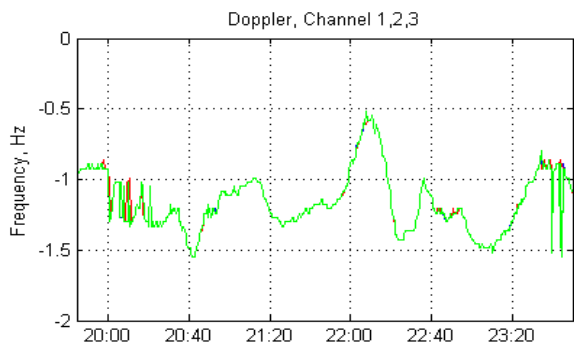
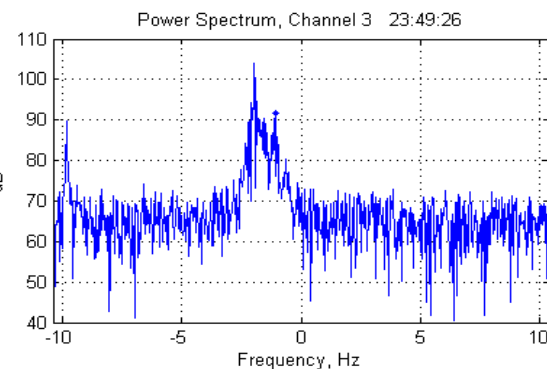
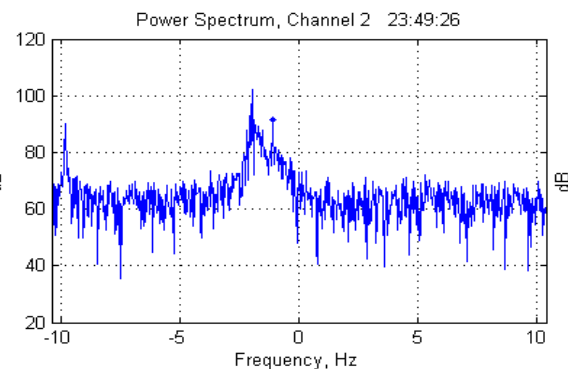
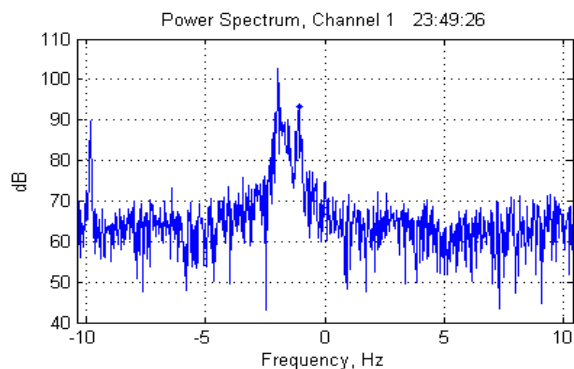
PC



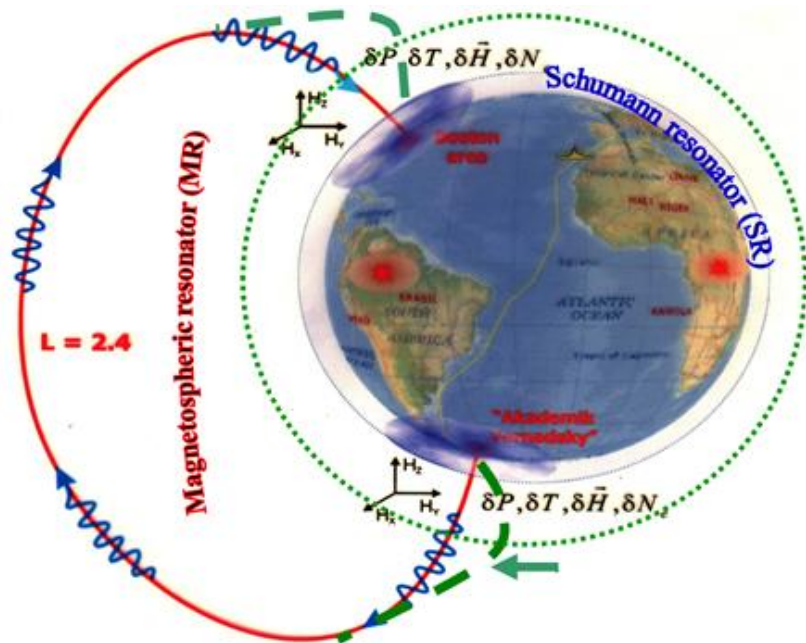
Rod antennas



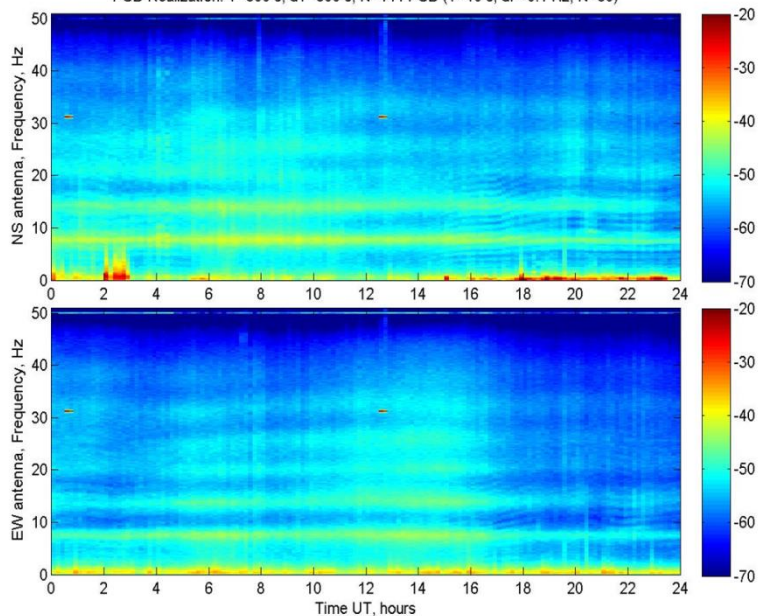
# Minsk – Kharkiv, D~700 km





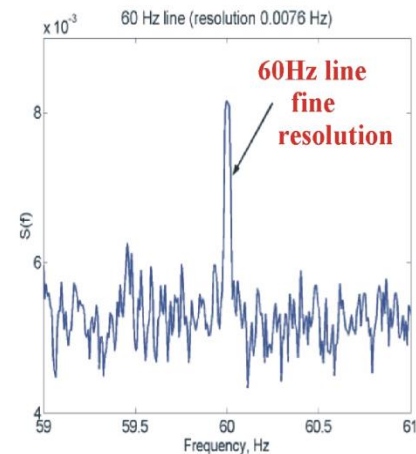
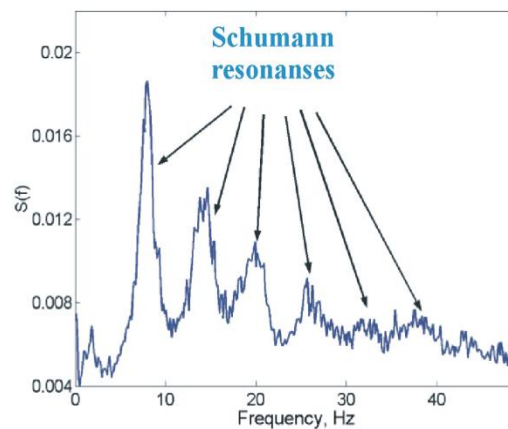
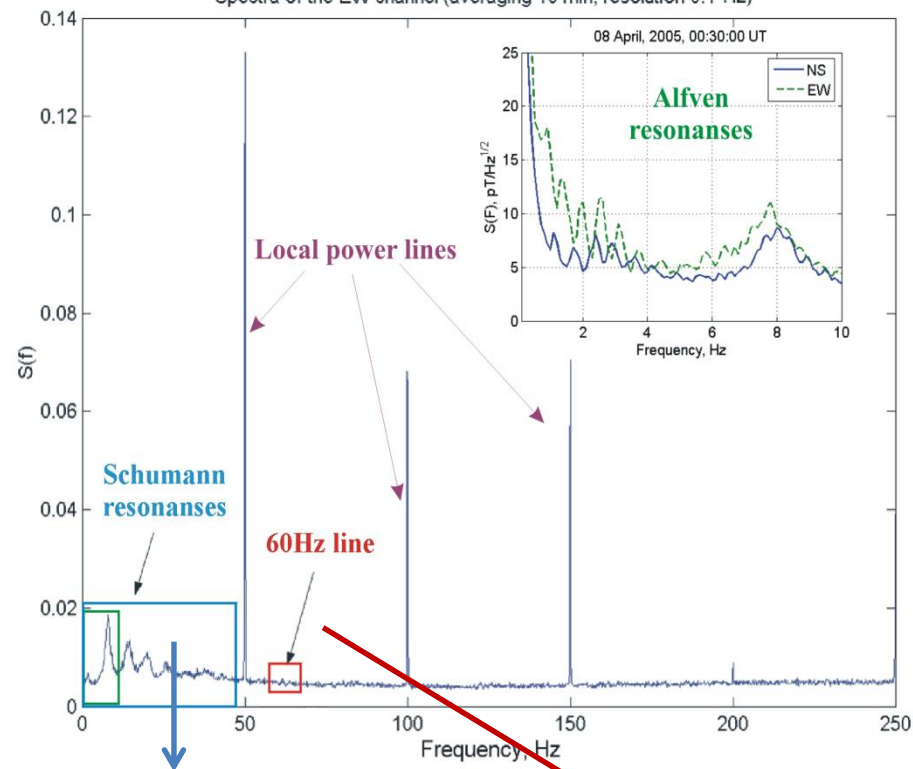


Specgram 01/06/2010 00:00:00-24:00:00 UT (Log Scale),  
PSD Realization: T=600 s, dT=600 s, N=144 PSD (T=10 s, dF=0.1 Hz, N=60)

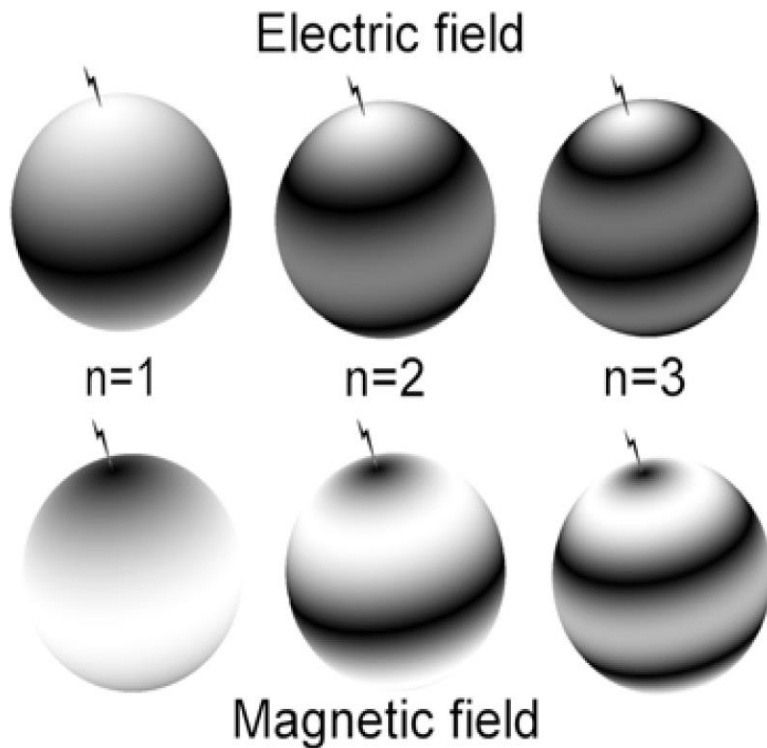


## Typical spectrum of ELF signal measured at Vernadsky

Spectra of the EW channel (averaging 10 min, resolution 0.1 Hz)

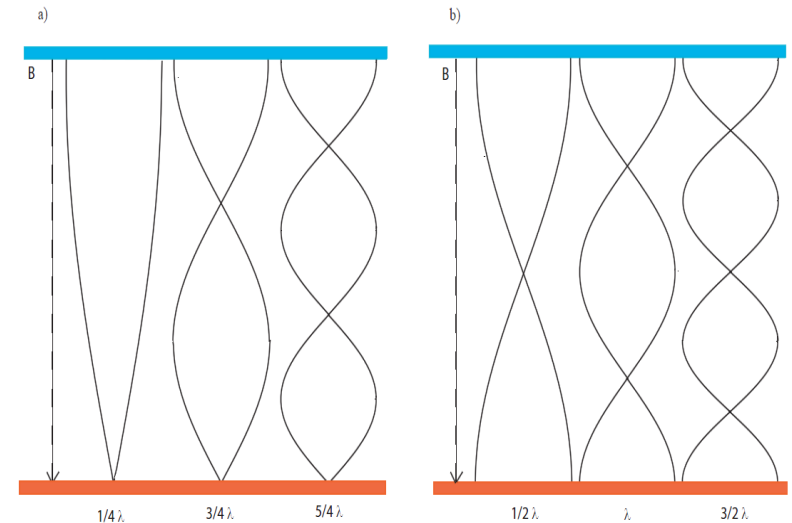


# SR



Electric and magnetic fields of the first three SR modes.

# IAR

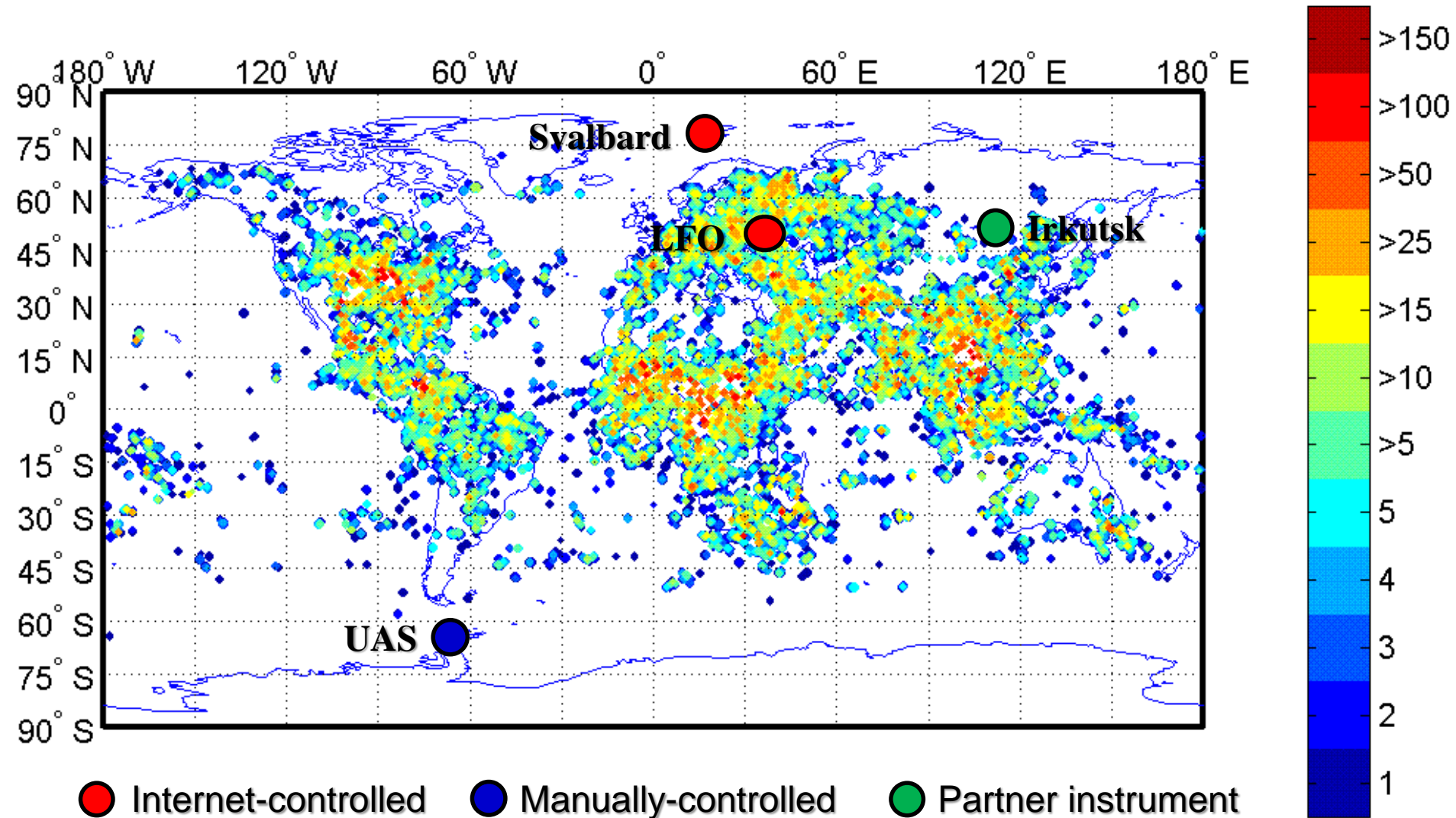


**Fig** Diagram showing an Alfvén wave's perpendicular electric field for the first 3 eigenmodes trapped in the IAR for the high conductivity ionosphere (panel (a)) and low conductivity (panel (b)) ionosphere. The lower red rectangles represent the lower IAR boundary and the blue rectangle represents the upper IAR boundary.

# Locations of the ELF waveband magnetometers and animation demonstrating distribution of lightning flashes (OTD satellite data)

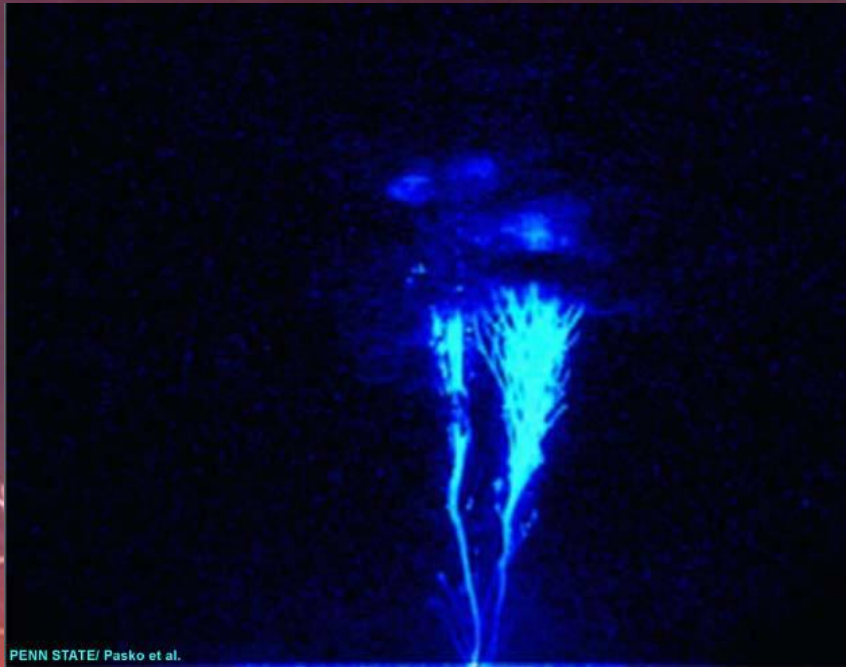
1995 spring

flashes





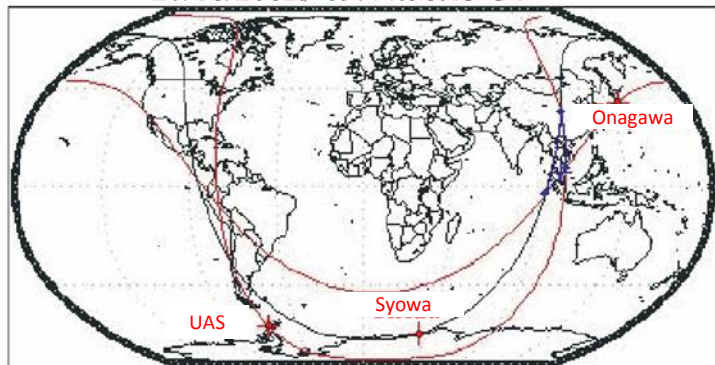
# Lightning is a main source of SR and IAR



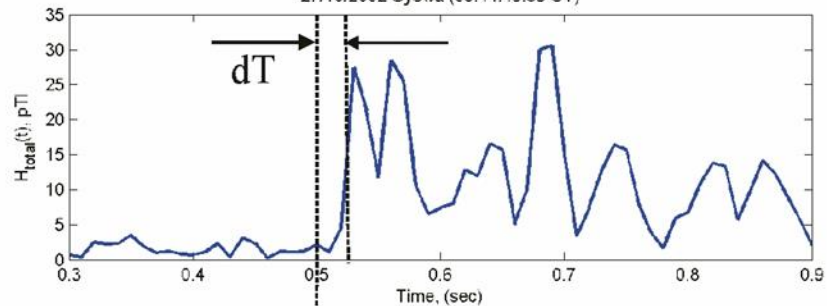
PENN STATE/ Pasko et al.



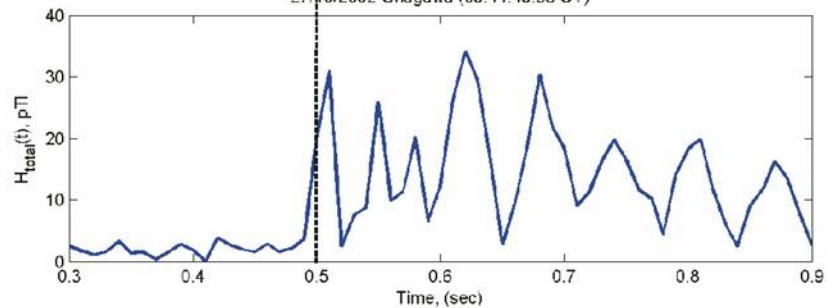
27/10/2002, 09:44:50.15 UT



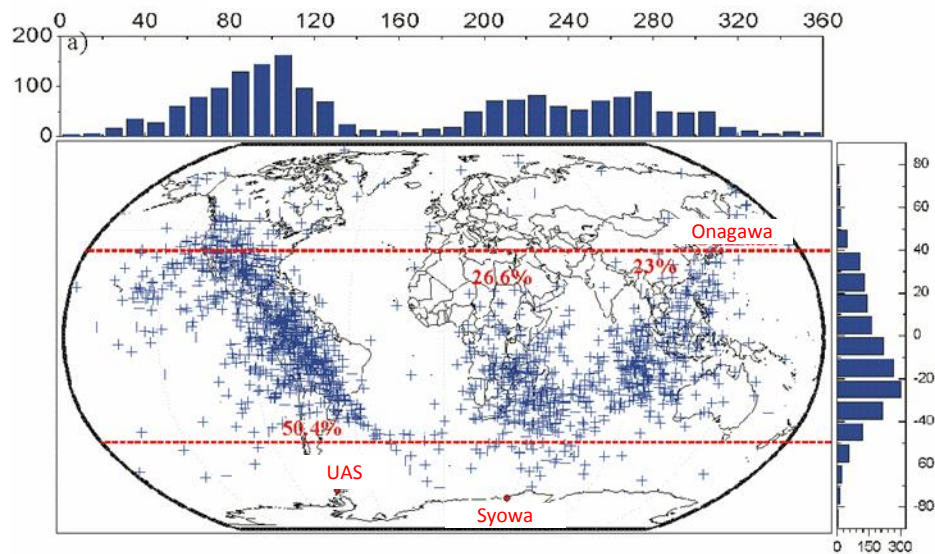
27/10/2002 Syowa (09:44:49.65 UT)



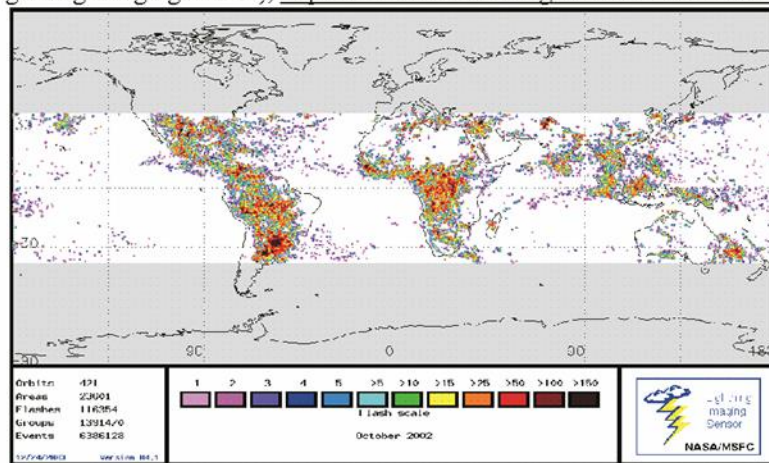
27/10/2002 Onagawa (09:44:49.65 UT)



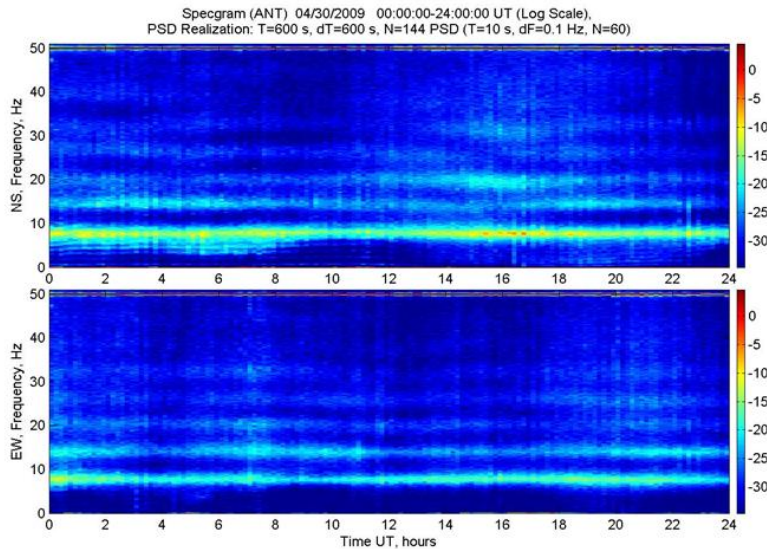
October 27-30, 2002



(Lightning Imaging Sensor), <http://thunder.msfc.nasa.gov/data/lisbrowse.html>.



# Spectral and polarization technique of data processing



Diurnal record is divided on intervals  $dT$ . For each interval power spectra  $S_{xx}$ ,  $S_{yy}$  and cross-spectrum  $S_{xy}$  are calculated

$$S_{ik}(f) = \frac{1}{N} \sum_{p=1}^N \tilde{H}_i^{(p)}(f) \tilde{H}_k^{(p)*}(f)$$

Instantaneous spectra  $\tilde{H}_{i,k}^{(p)}(f)$  corresponding to the intervals  $(p-1)T \leq t \leq pT$ ;  $p = 1, 2, \dots, N$ ,  $N \cdot T = dT$  are determined by

$$\tilde{H}_{i,k}^{(p)}(f) = \frac{1}{T} \int_{(p-1)T}^{pT} dt H_{i,k}(t) e^{-i2\pi ft}$$

Stockes parameters are calculated

$$I(f) \equiv S_{xx}(f) + S_{yy}(f), \quad Q(f) \equiv S_{xx}(f) - S_{yy}(f),$$

$$U(f) \equiv 2 \operatorname{Re} S_{xy}(f), \quad V(f) \equiv 2 \operatorname{Im} S_{xy}(f)$$

and used for determination of the:

$r(f)$  - Ellipticity

$\Psi(f)$  - Polarization angle

$I_p(f)$  - Polarized intensity

$P(f)$  - Degree of polarization

$$r = \frac{V}{\sqrt{Q^2 + U^2} + \sqrt{Q^2 + U^2 + V^2}},$$

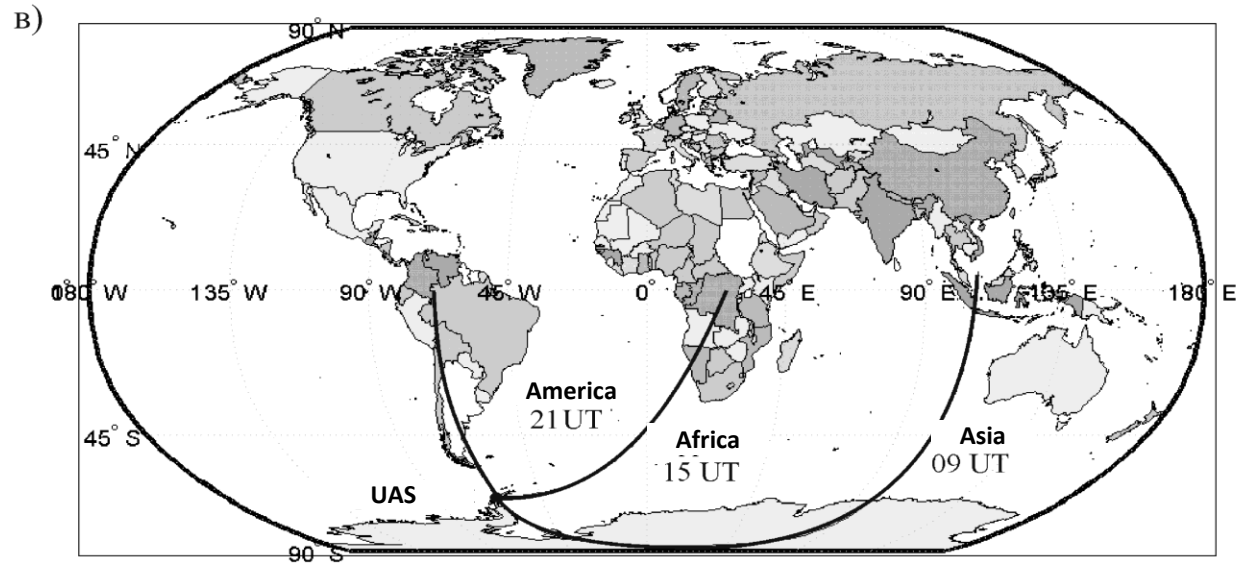
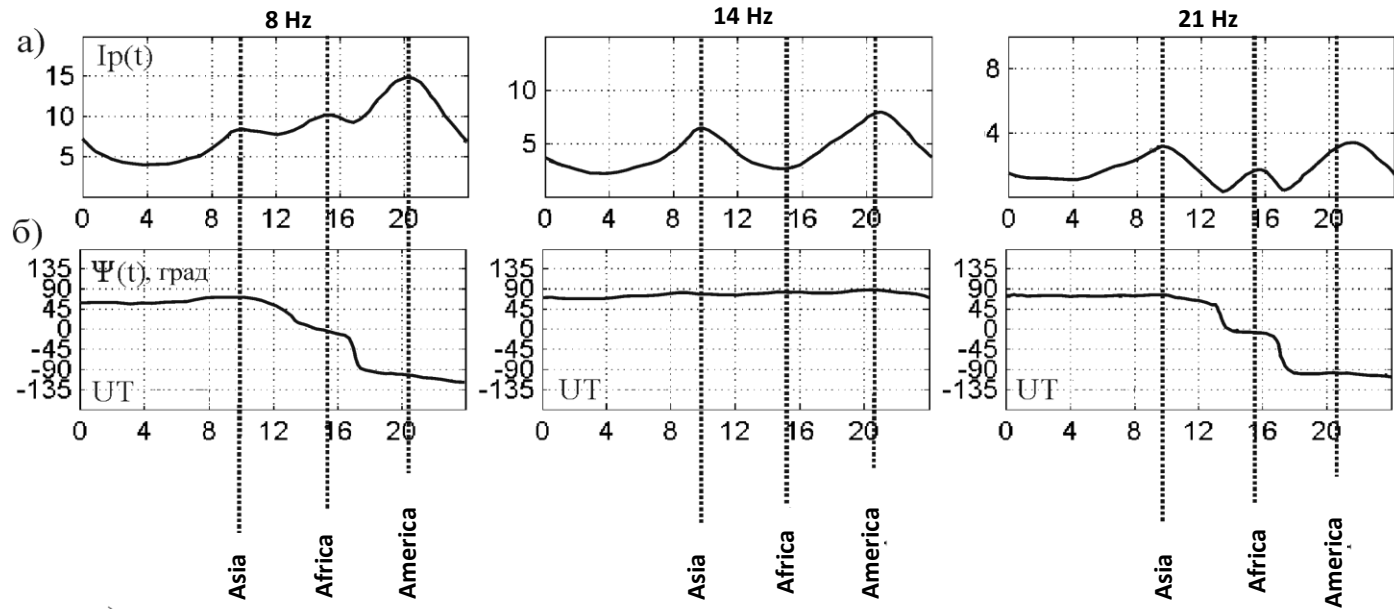
$$\left. \begin{matrix} \sin(2\Psi) \\ \cos(2\Psi) \end{matrix} \right\} = \frac{1}{\sqrt{Q^2 + U^2}} \left\{ \begin{matrix} U \\ Q \end{matrix} \right\},$$

$$I_p = \sqrt{Q^2 + U^2 + V^2},$$

$$P = I_p / I.$$

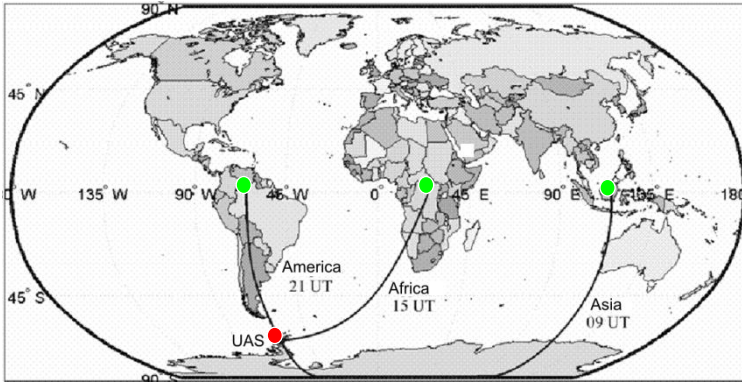


# Intensities and position angle of SR signals



# Calculation of the activities of the world lightning centers

In the frame of asymptotic theory of fields of the Schumann resonance in the gyrotropic Earth-ionosphere cavity



- Lightning activity is concentrated in 3 points centers at the equator
- The Earth surface is perfectly conducted
- Gyrotropic ionosphere with high-conductivity boundary is described by the tensor the of the surface impedance  $Z$  with small diagonal element  $|Z_{ik}| \ll 1$
- Activities of Asia –  $A_1$  and Africa+America –  $A_2+A_3$  lightning centers are calculated

Impedance of the isotropic ionosphere  $Z_0$  and the ratio of the off-diagonal and diagonal elements of the matrix of the surface impedance of the gyrotropic ionosphere  $Z_2/Z_1$  are used as small parameters

$$A_1^{(1)}(t) = \frac{1}{2} [A_+^{(1)}(t) + A_-^{(1)}(t)] \quad - \text{Africa}$$

$$[A_2^{(1)}(t) + A_3^{(1)}(t)] = \frac{1}{2} [A_+^{(1)}(t) - A_-^{(1)}(t)] \quad - \text{Asia + America}$$

$$\rho \equiv \frac{\omega_H}{\nu_e} = \frac{1.5 \cdot \Delta\lambda^{(1)}}{(0.75)^2 - (\Delta\lambda^{(1)})^2} \quad - \text{Parameter of gyrotropy}$$

$$\text{rде: } A_+^{(1)}(t) = \frac{1}{2 \sin^2 \Phi} \{ I_p(f^{(1)}, t) (1 + \sin^2 \Phi) - Q(f^{(1)}, t) \cos^2 \Phi \}$$

$$A_-^{(1)}(t) = \{ I_p(f^{(1)}, t) \cos^2 \Phi - Q(f^{(1)}, t) (1 + \sin^2 \Phi) \} \begin{cases} 1/2 \sin^2 \Phi \sin 2(\lambda - \lambda_1^{(1)}) - \text{gyrotropic cavity} \\ 1/2 \sin^2 \Phi \sin 2(\lambda - \lambda_1) - \text{isotropic cavity} \end{cases}$$

$$\Delta\lambda^{(1)} = -\frac{V(f^{(1)}, t)}{2A_+^{(1)} \sin \Phi} \quad - \text{«apparent» angle shift of Africa source to the west } \hat{\lambda}_1^{(1)} = \lambda_1 - \Delta\lambda^{(1)}$$

$A_j^{(1)}(t)$  - electromagnetic activity (j-source) for first Schumann frequency

$\{\Phi; \lambda\}$  - observation point coordinates  $\{\Phi_j = 0; \lambda_j\}$  - lightning center coordinates.

$$A_j(t) \equiv \left[ \frac{3c}{4\pi^2 a^3 f^{(1)} |z_1|} \right]^2 < |\tilde{M}_j(t)|^2 > \quad - \text{electromagnetic activity (j-source), where } a - \text{Earth radius, } c - \text{speed of light,}$$

$f^{(1)}$  - frequency of first maximum in the point of source

$< |\tilde{M}_j(t)|^2 >$  - mean square of the current moment of j-discharge

$$z_1 = \frac{z_0}{\sqrt{2}} \sqrt{1 + \sqrt{1 + \rho^2}} \quad - \text{diagonal element of the impedance}$$

$$z_0 = \sqrt{\frac{i\omega\nu_{eff}}{\omega_0^2}} \quad - \text{impedance of the isotropic ionosphere}$$

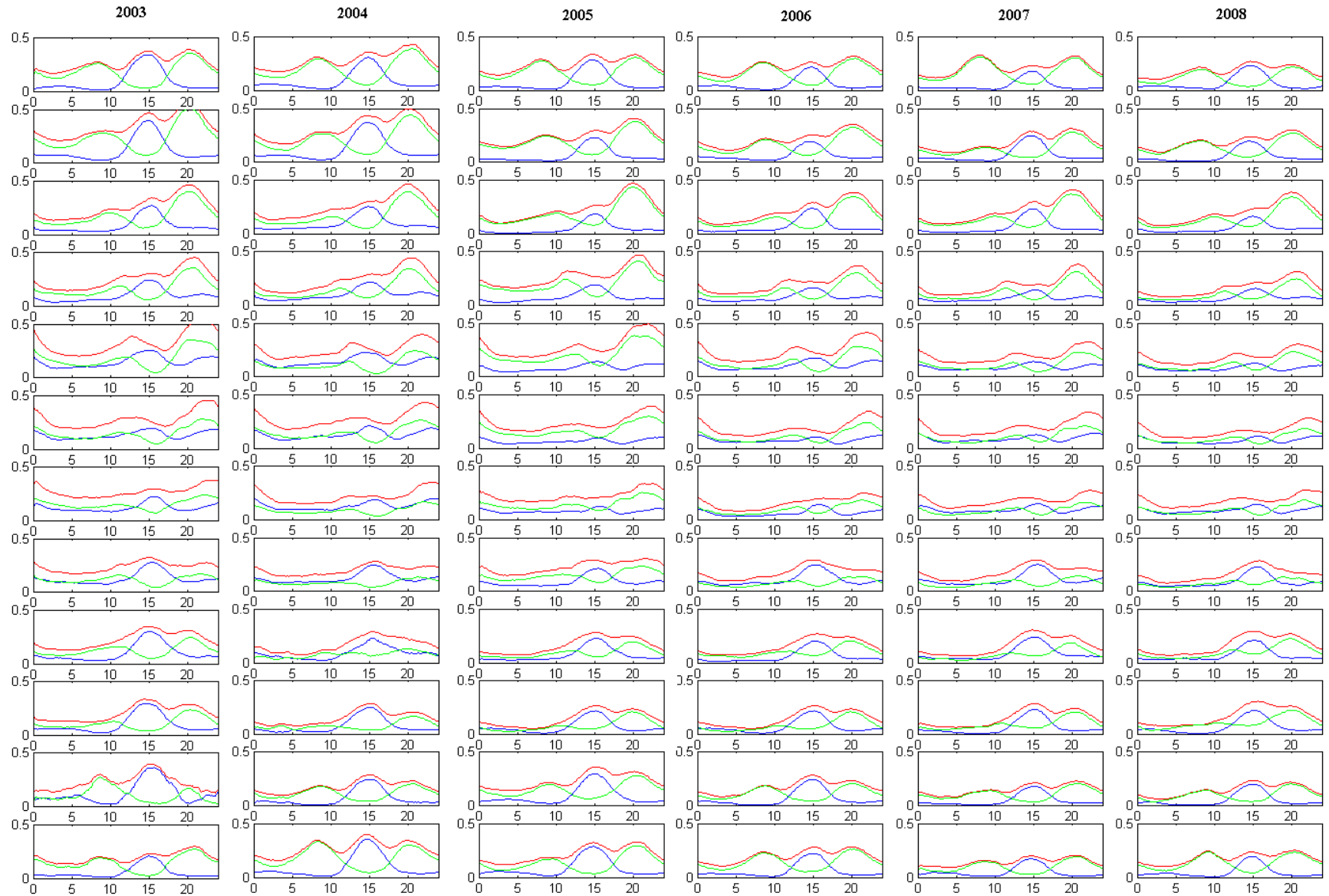
$\omega_0$  - critical frequency of the ionosphere

$\omega_H$  - ion gyro frequency

$\nu_{eff}$  - electron collision frequency

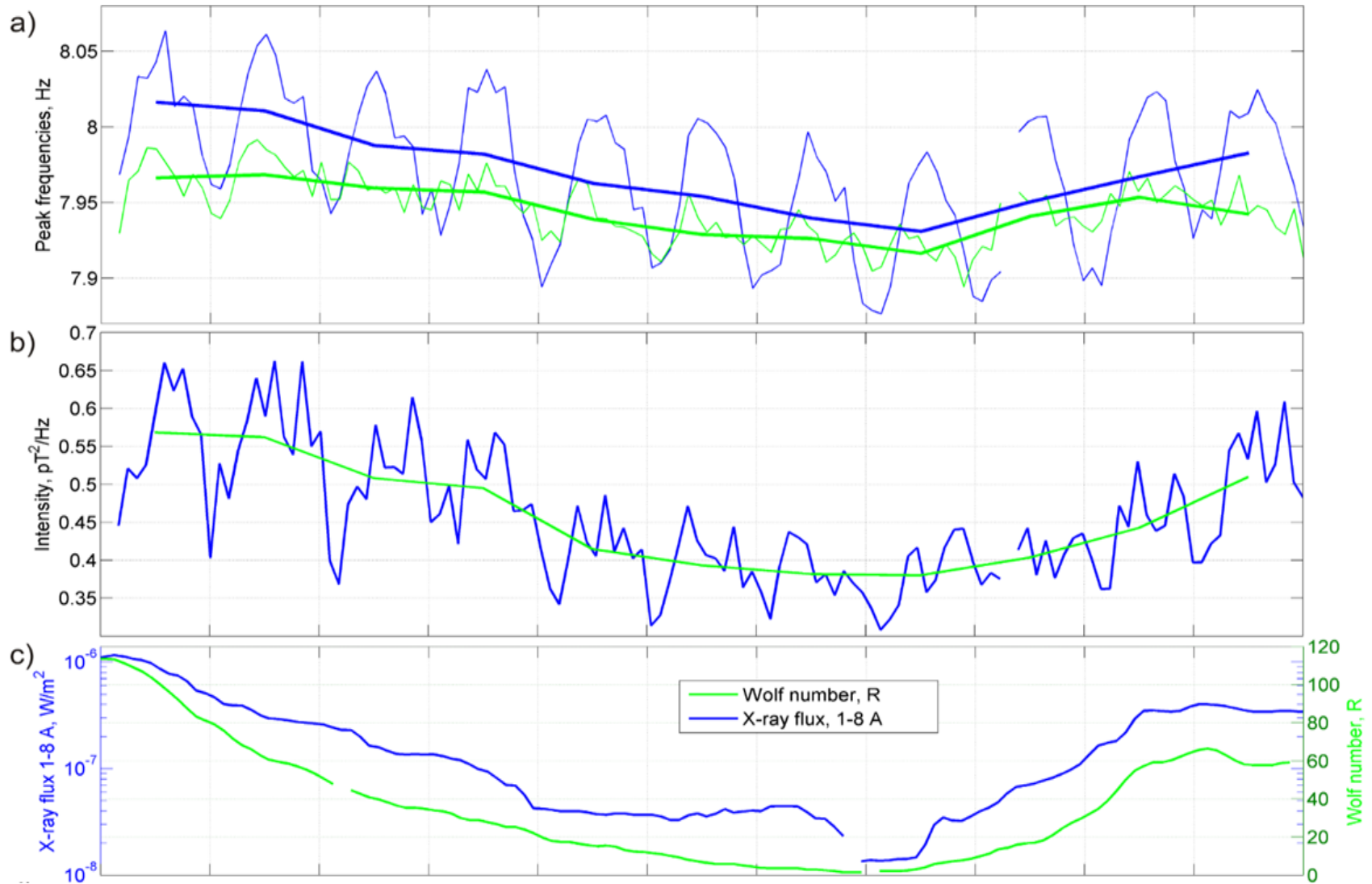
# Diurnal and seasonal variation of the activities of world lightning centers

( $A_1 + A_{23}$  – Total activity,  $A_1$  – Africa,  $A_{23}$  – Asia + America)



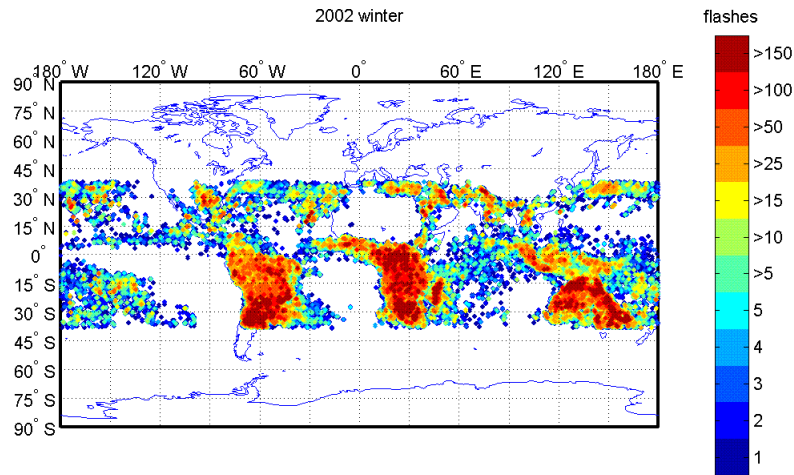


# Parameters calculated within 11-years solar cycle

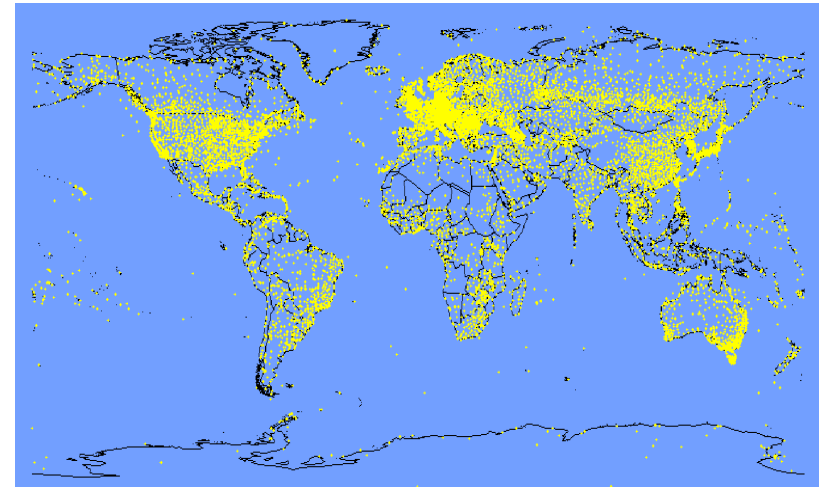


# Estimation of the characteristics of Schumann resonance signal sources

## Distributions of the numbers of flashes

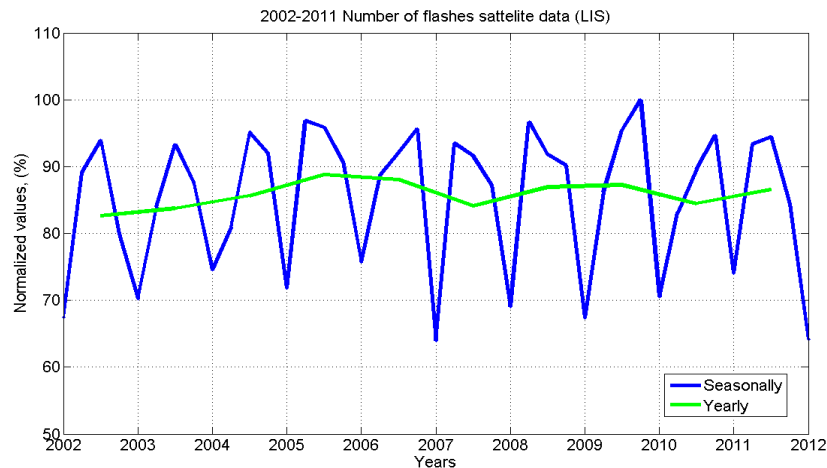


## Map of the meteorological stations



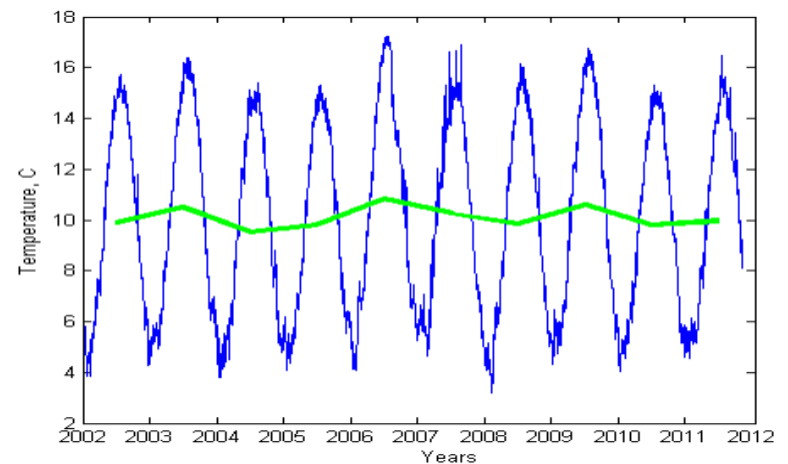
## Variation of the number of flashes calculated with satellite data (Lightning Imaging Sensor (LIS) data:

<https://amser.org/index.php?P=AMSER--ResourceFrame&resourceId=4353>



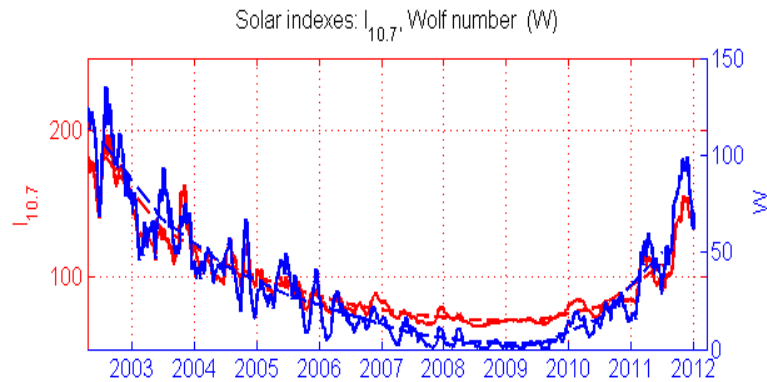
## Variations of daily average temperature on the surface of the Earth from 2002 to 2012 (blue line) its average value (green line)

<ftp://ftp.ncdc.noaa.gov/pub/data/gsod>

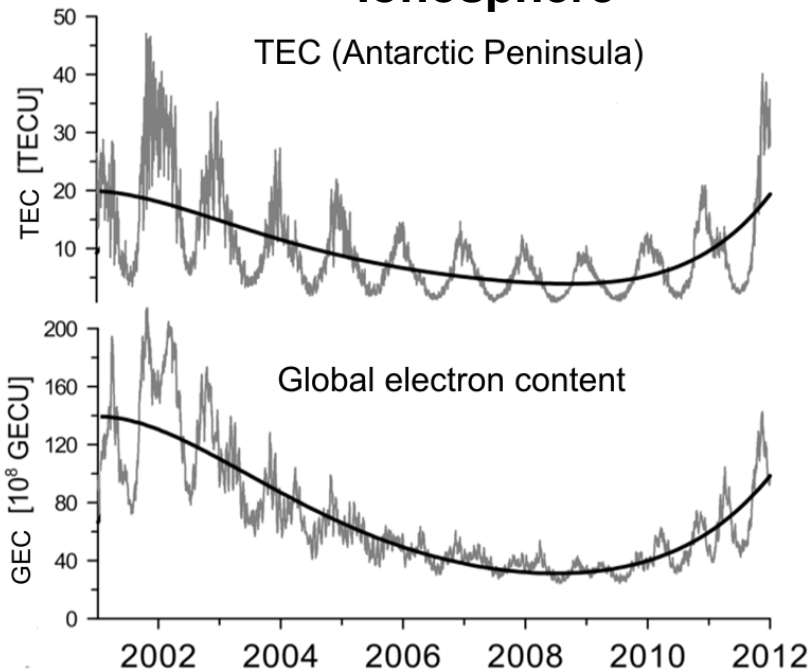


# Long-term trends in Sun/Ionosphere parameters

## Sun



## Ionosphere



## The ways to improve the model

$$A_j(t) \equiv \left[ \frac{3c}{4\pi^2 a^3 f^{(1)} |z_1|} \right]^2 < |\tilde{M}_j(t)|^2 > - \text{electromagnetic activity (j-}$$

source), where  $a$  - Earth radius,  $c$  - speed of light,

$f^{(1)}$  - frequency of first maximum in the point of source

$< |\tilde{M}_j(t)|^2 >$  - mean square of the current moment of j-discharge

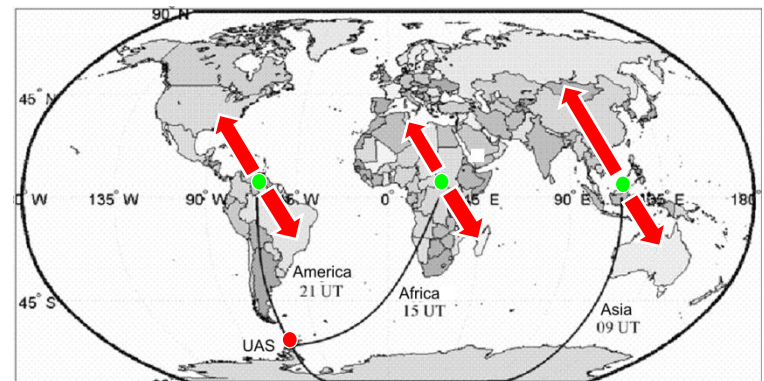
$$z_1 = \frac{z_0}{\sqrt{2}} \sqrt{1 + \sqrt{1 + \rho^2}} - \text{diagonal element of the impedance}$$

$$z_0 = \sqrt{\frac{i\omega\nu_{eff}}{\omega_0^2}} - \text{impedance of the isotropic ionosphere}$$

$\omega_0$  - critical frequency of the ionosphere

$\omega_H$  - ion gyro frequency

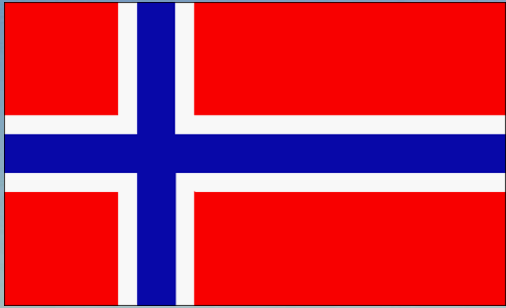
$\nu_{eff}$  - electron collision frequency





# Summary

- The facilities are integrated into **single Internet based system** which provides **remote control** of the instruments and **transferring of the raw data** to main server for further **processing** and **visualizing on the web page**. This scheme may be effectively used for organizing practical training courses for the students.
- Compact-sized **HF** and **ELF facilities** may be used both for studying of the process in the near Earth environment and **as educational tools** for **visualization** of these processes and **calculation** of their characteristics.



**Thank you!**